

Modeling Cl₂ formation from aqueous NaCl particles: Evidence for interfacial reactions and importance of Cl₂ decomposition in alkaline solution

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[1] A series of experiments have demonstrated that a significant yield of chlorine gas is produced when mixtures of ozone and sodium chloride particles above their deliquescence point are irradiated at 254 nm. In order to obtain expressions for future modeling studies, a comprehensive model is used to analyze the system and to determine its sensitivity. This work reexamines and expands previous studies [Knipping *et al.*, 2000]. The enhanced model, described in detail herein, reaffirms that current known physical and chemical processes fail to reproduce the observed Cl₂ formation in the experiments. A methodological analysis, proposed as a framework for similar studies, of the physicochemical system supports the accountability of an overall mechanism initiated by the formation of a relatively stable complex of the hydroxyl radical and chloride ions at the gas–liquid interface for the observed chlorine generation. Different potential fates of the OH•••Cl_{surface}[−] intermediate are discussed. A rate expression and kinetic parameters are presented for the overall reaction of the interfacial mechanism. In addition, sensitivity studies underscore the importance of accurately modeling chlorine decomposition processes in alkaline solution—in particular, the reactions of chlorine with hydroxide, carbonate, and basic hydrogen peroxide. Recommended aqueous-phase rate constants for these reactions are drawn from a literature evaluation illustrating the limited availability and lack of agreement of related kinetic data. **INDEX TERMS:** 0305 Atmospheric Composition and Structure: Aerosols and particles (0345, 4801); 0312 Atmospheric Composition and Structure: Air/sea constituent fluxes (3339, 4504); 0317 Atmospheric Composition and Structure: Chemical kinetic and photochemical properties; 0365 Atmospheric Composition and Structure: Troposphere—composition and chemistry

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1. Introduction

[2] The role of chlorine atoms as potential oxidants in the marine boundary layer and coastal urban areas has received a great deal of attention in recent years [Keene *et al.*, 1990, 1996, 1998; Finlayson-Pitts, 1993; Pszenny *et al.*, 1993; Finlayson-Pitts and Pitts, 2000; Finlayson-Pitts and Hemminger, 2001]. Laboratory and modeling studies have shown that photolabile chlorine-containing species are generated from several reactions involving sea salt particles [Vogt *et al.*, 1996; Sander and Crutzen, 1996; Langer *et al.*, 1997; De Haan *et al.*, 1999; Weis and Ewing, 1999; Massucci *et al.*, 1999; Regimbal and Mozurkewich, 2000]. In addition, intriguing evidence from field and laboratory studies indicates that aqueous sea salt particles serve as a potential source of molecular chlorine, Cl₂ [Spicer *et al.*, 1998; Oum *et al.*, 1998].

[3] Spicer *et al.* [1998] have detected concentrations of molecular chlorine in the marine boundary layer at levels

higher than those expected from known reactions of sea salt particles. Using an atmospheric pressure chemical ionization mass spectrometer, Spicer *et al.* measured Cl₂ mixing ratios as high as 150 parts per trillion (ppt) during nighttime at a site off the mid-Atlantic coast of the United States. Using a photochemical box model, the investigators concluded that the observed Cl₂ levels could not be reconciled by known chemistry and inferred the presence of an unrecognized continuous chlorine source averaging up to 330 Cl₂ ppt day^{−1}. Whereas the peak Cl₂ concentrations measured by Spicer *et al.* are indicative of a nighttime mechanism for Cl₂ build-up, Oum *et al.* [1998] present evidence that could explain a daytime component for Cl₂ release from deliquesced sea salt aerosol.

[4] Oum *et al.* [1998] performed laboratory experiments in a modern aerosol chamber that identify Cl₂ as the product of the photolysis of ozone in the presence of deliquesced sea salt particles. The use of both synthetic sea salt and sodium chloride particles resulted in similar rates of Cl₂ production. The experiments are able to generate Cl₂ only when the particles are above their deliquescence point and the chamber contents are photolyzed; Cl₂ production is not observed

either when the deliquesced particles and ozone remain in the dark or when the chamber contents are irradiated in the absence of ozone. Oum et al. report a small production of Cl_2 when "dry" sea salt particles are introduced into the chamber, attributed to reactions occurring on films of adsorbed water known to persevere on the surface of the particles. Additional experiments were performed in a clean 70 L Teflon reaction chamber in order to ensure that the observed Cl_2 production was not a consequence of some unknown chamber contaminant or artifact. Cl_2 formation was observed in conditions similar to those of the aerosol chamber.

[5] Oum et al. proposed a mechanism involving $\text{OH}^\bullet$ oxidation of the chloride ion, Cl^- , as the potential pathway for the observed Cl_2 production. Although this mechanism includes paths of chlorine production based on known bulk aqueous-phase chemistry, model simulations will show that it fails to reproduce observed Cl_2 values by as much as four orders of magnitude.

[6] There is increasing recognition that aqueous aerosol chemistry is not limited to bulk liquid-phase reactions. Laboratory studies have shown that Cl_2 reacts with bromide ions at aqueous aerosol interfaces, forming a surface complex species $(\text{Cl}_2 \bullet \text{Br})^-$ and leading to enhanced uptake of the gas [Hu et al., 1995]. Additional experiments uphold similar conclusions for reactions of SO_2 with water at droplet surfaces [Jayne et al., 1990; Boniface et al., 2000]. Second harmonic generation and static surface tension measurements of SO_2 surface coverage support the hypothesis of SO_2 adsorption on liquid surfaces [Donaldson et al., 1995].

[7] There is strong evidence suggesting that the underlying mechanism for the observed Cl_2 production in the aerosol chamber experiments may involve similar reactions occurring at the interface of deliquesced sodium chloride particles. Molecular dynamics simulations illustrate that chloride ions are much more prevalent than sodium ions at the surfaces of concentrated NaCl solutions. Furthermore, quantum chemical calculations predict that these chloride ions will attract reactive gases such as $\text{OH}^\bullet$ and O_3 [Knipping et al., 2000; Jungwirth and Tobias, 2000]. The probability of these reactions is explored in this research.

[8] Clearly, the formation of molecular chlorine from the reaction of ozone with aqueous sea salt particles involves the simultaneous occurrence of a number of interrelated chemical and physical processes. In order to analyze the experimental results and explore the viability of different chemical reaction mechanisms, a new mathematical model simulating the physical and chemical processes within the experimental chamber used by Oum et al. [1998] is presented. The model incorporates a comprehensive mechanism of gas-phase and aqueous-phase chemical reactions in conjunction with a dynamic treatment of gas-phase to aqueous-phase mass transport. Furthermore, the model also considers reactions occurring at the gas-particle interface. Hence, the new model is named Model of Aqueous, Gaseous and Interfacial Chemistry, or MAGIC. Details of the model are presented herein, including the latest revisions.

[9] The results from MAGIC are analyzed in order to test various pathways for the observed Cl_2 formation. Once a suitable rationalization for the phenomena is reached, an appropriate mechanism and overall rate expression are developed. This mechanism should be subjected to further

appraisal by complex atmospheric models such as MOCCA [Sander and Crutzen, 1996], CAPRAM [Herrmann et al., 2000] and future expanded versions of MAGIC. As chlorine atoms, resulting from the subsequent photolysis of Cl_2 , can affect the chemistry of the marine boundary layer and coastal regions, this mechanism will assist models in better determining the significance of chlorine chemistry in these regions. An added goal of this research is to appeal to the scientific community to further investigate alkaline chlorine decomposition processes found essential to the analysis of the physicochemical system at hand through a suite of sensitivity analyses.

2. Model Description

[10] The model developed for the current study simulates the conditions present in the aerosol chamber. A description of the aerosol chamber and experiments is found elsewhere [De Haan et al., 1999; Knipping et al., 2000]. Briefly, sodium chloride aerosol is introduced into a stainless-steel chamber, whose walls are coated with halocarbon wax to render them inert, at a relative humidity above its deliquescence point and at ambient temperature. Ozone is introduced into the chamber and Cl_2 production is measured using an atmospheric pressure ionization mass spectrometer after photolysis of the chamber contents.

[11] The model contains a comprehensive treatment of gas-phase and aqueous-phase chemical reactions, using the most recent kinetic data available, as well as mass transfer of species between the gas phase and the aqueous phase. We refer the reader to the seminal works of Fuchs and Sutugin [1971] and Schwartz [1986], as well as a recent review by Sander [1999], for a detailed overview of the modeling of gas-phase and aqueous-phase interactions.

[12] The steps involved in the dissolution, and subsequent aqueous-phase chemical reaction, of gas-phase species into droplets are well established in the literature [Schwartz and Freiberg, 1981; Schwartz, 1986; Kumar, 1989; Lelieveld and Crutzen, 1991]. These processes occur in parallel with chemical reaction in the gas phase and are summarized as follows:

1. Transport of the species from the gas phase toward the droplet gas-liquid interface.
2. Transfer of the species across the interface.
3. Possible hydrolysis and partial ionization of the species in the aqueous phase.
4. Diffusion of the ionic and nonionic species in the droplet.
5. Aqueous-phase chemical reaction.

[13] Steps 1–5 may also proceed in reverse order, potentially generating concentration gradients at the droplet interface that lead to the volatilization of species to the gas phase. In addition, MAGIC also considers reactions on the surface of the aerosol particles. These reactions are also assumed to occur in parallel with the physical and chemical processes described above.

2.1. Gas-Phase Chemistry

[14] The treatment of gas-phase chemical reactions in the model adheres to the recommendations made by the NASA/JPL evaluation panel [DeMore et al., 1997; Sander et al., 2000] which are in general agreement with the recommen-

Table 1. List of Species Included in the Model for Laboratory Simulations

Group	Gas phase	Aqueous phase
HO _x	O(¹ D), O(³ P), O ₂ , O ₃ , OH [•] , HO ₂ [•] , H ₂ O ₂ , H ₂ O	O(³ P), O ₂ , O ₃ , OH [•] /O ^{•−} , HO ₂ [•] /O ₂ ^{•−} , H ₂ O ₂ /HO ₂ [−] , HO ₃ [•] , O ₃ ^{•−} , H ⁺ /OH [−]
ClO _x	Cl ₂ , Cl [•] , HCl, HOCl, Cl ₂ O, ClO [•] , OClO [•] , ClOO [•] , Cl ₂ O ₂	Cl ₂ , Cl [•] , Cl ₂ ^{•−} , HCl/Cl [−] , HOCl/ClO [−] , Cl ₂ O, HOClH [•] , HOCl ^{•−}
CO ₂	CO ₂	CO ₂ •H ₂ O/HCO ₃ [−] /CO ₃ ^{2−} , HCO ₃ [•] /CO ₃ ^{•−}
Other	N ₂	Na ⁺

Species in bold are transferred between the aqueous phase and the gas phase.

dations made by the IUPAC panel [Atkinson *et al.*, 1997, 2000]. Table 1 lists the species considered in the gas and aqueous phase during these simulations.

[15] Low-pressure mercury lamps used in the experiments were chosen because their primary emission line, at a wavelength, λ , of 254 nm, photolyzes ozone but not Cl₂. The photolysis rates, j , of the gas-phase species (see Table 2) are evaluated by first experimentally determining the photolysis rate of ozone in the aerosol chamber. The photolysis rates for all other species are then computed by taking the ratio of the absorption cross section of the species with respect to the absorption cross section of ozone at $\lambda = 254$ nm and multiplying the ozone photolysis rate by this value. The relative quantum yields at 254 nm are also considered. Some Cl₂ is observed to photolyze due to weaker emission lines (e.g., 313 and 365 nm). The rate of the resulting Cl₂ photolysis in the chamber is quantified experimentally and included in the model. The kinetic data for gas-phase chemical reactions of interest for the chamber studies are presented in Table 3.

2.2. Activity Coefficients

[16] Long-range and short-range electrostatic forces, or interactions, between ions affect the reactivity of species in concentrated ionic solutions. The ionic strength, I , represents the appropriate measure of the total ion concentration of a solution taking into account these effects,

$$I = \frac{1}{2} \sum_i m_i z_i^2 \quad (1)$$

where m and z are the molality and valence of species i , respectively [Debye and Hückel, 1923].

[17] The activity of a solute and its concentration are related by activity coefficients, γ . Several methods exist for calculating activity coefficients in ionic solutions [Kim *et al.*, 1993]. The 1991 Pitzer ion interaction model [Pitzer, 1991] is used in this study. The Pitzer interaction model has been validated to calculate activity coefficients in multi-component solutions with ionic strengths up to 6 mol kg^{−1}. This limit may restrict the modeling of highly concentrated atmospheric aerosol but it does not hinder its applicability in this study as this value is above that for saturation of NaCl in water. The Pitzer model can also be expanded to calculate the activity coefficients of neutral solutes, such as CO₂ and O₂, in ionic solutions [Clegg and Brimblecombe, 1990].

[18] Pitzer ion interaction data are available for a limited number of possible solutes, representing the salts most prevalent in atmospheric aerosol. The model calculates activity coefficients explicitly for the species H⁺, Na⁺, Cl[−], HCO₃[−], CO₃^{2−}, OH[−], CO₂, and O₂. For other species, either a Guggenheim approximation of the Debye–Hückel limiting law [Pethybridge and Prue, 1972],

$$-\log \gamma_i = \frac{Az_i^2 I^{1/2}}{1 + I^{1/2}}, \quad (2)$$

is employed, where A is the Debye–Hückel constant for activity coefficient, 0.509 (kg mol^{−1})^{1/2} at 298 K, or, in the case of neutral solutes, activity coefficients are calculated using the approximation [Moldanova and Ljungstrom, 2001]:

$$\log \gamma_i = 0.1I. \quad (3)$$

2.3. Mass Transfer

[19] In order to quantify the limitations to mass transport imposed by these physical processes, the mass transfer coefficient, k_{mt} , developed by Schwartz [1986] is employed. The mass transfer coefficient is effectively a first-order rate constant that determines the steady state rate by which a gas enters (or escapes) a droplet in response to (spatially averaged) chemical reaction loss and production within the bulk solution, $\langle R_{aq} \rangle$, namely,

$$k_{mt} \left(C_{g,\infty} - \frac{C_{aq,s}}{H^*RT} \right) = \langle R_{aq} \rangle \quad (4)$$

where

$$k_{mt} = \left[\frac{R_p^2}{3D_g} + \frac{4R_p}{3\bar{c}\alpha} \right]^{-1} \quad (5)$$

and $C_{g,\infty}$ and $C_{aq,s}$ represent the bulk gas-phase and surface aqueous-phase concentration of the species of interest, H^* is the effective Henry's law constant, R is the universal gas constant and T denotes absolute temperature. The droplet radius is denoted by R_p , while D_g , $\bar{c}$ and α represent the gas-phase diffusivity, mean molecular speed and accommodation coefficient for the species of interest, respectively. The value of droplet radius is found using a quasi-size-dependent approach [Sander and Crutzen, 1996]. Mathematically, k_{mt} represents the sum in series of the resistances to uptake due to gas-phase diffusion and interfacial mass transport.

[20] Although Henry's law equilibrium is not imposed in the model, the rate of mass transfer to and from the gas

Table 2. Gas-Phase Photolysis Rates

No.	Reaction	σ ($\lambda = 254$ nm) $\times 10^{20}$ cm ²	Φ	j_{298} s ^{−1}
1	O ₃ + $h\nu \rightarrow$ O(³ P) + O ₂	1150	0.1	4.50×10^{-4}
2	O ₃ + $h\nu \rightarrow$ O(¹ D) + O ₂	1150	0.9	4.05×10^{-3}
3	H ₂ O ₂ + $h\nu \rightarrow$ OH [•] + OH [•]	6.7	1.0	2.62×10^{-5}
4	HOCl + $h\nu \rightarrow$ OH [•] + Cl [•]	14.6	1.0	5.71×10^{-5}
5	Cl ₂ + $h\nu \rightarrow$ Cl [•] + Cl [•]	0.0	—	1.0×10^{-4a}
6	OClO [•] + $h\nu \rightarrow$ ClO [•] + O(³ P)	0.0	—	0
7	Cl ₂ O ₂ + $h\nu \rightarrow$ ClOO [•] + Cl [•]	544.5	1.0	2.13×10^{-3}
8	Cl ₂ O + $h\nu \rightarrow$ Cl [•] + ClO [•]	193	1.0 ^b	7.55×10^{-4}

Absorption cross sections: DeMore *et al.* [1997].

^a Measured in chamber.

^b Estimated.

Table 3. Gas-Phase Chemical Mechanism

No.	Reaction	Order	<i>A</i>	<i>E/R</i>	<i>k</i> ₂₉₈ ²
9	O(¹ D) + O ₂ → O(³ P) + O ₂	2	3.2 × 10 ⁻¹¹	-70	4.0 × 10 ⁻¹¹
10	O(¹ D) + O ₃ → O ₂ + O ₂	2	1.2 × 10 ⁻¹⁰	0	1.2 × 10 ⁻¹⁰
11	O(¹ D) + O ₃ → O(³ P) + O(³ P) + O ₂	2	1.2 × 10 ⁻¹⁰	0	1.2 × 10 ⁻¹⁰
12	O(¹ D) + H ₂ O → OH• + OH•	2	2.2 × 10 ⁻¹⁰	0	2.2 × 10 ⁻¹⁰
13	O(¹ D) + N ₂ → O(³ P) + N ₂	2	1.8 × 10 ⁻¹¹	-110	2.6 × 10 ⁻¹¹
14	O(³ P) + O ₂ + M → O ₃ + M	2			1.5 × 10 ⁻¹⁴
15	O(³ P) + O ₃ → O ₂ + O ₂	2	8.0 × 10 ⁻¹²	2060	8.0 × 10 ⁻¹⁵
16	O(³ P) + HO ₂ • → OH• + O ₂	2	3.0 × 10 ⁻¹¹	-200	5.9 × 10 ⁻¹¹
17	O(³ P) + H ₂ O ₂ → OH• + HO ₂ •	2	1.4 × 10 ⁻¹²	2000	1.7 × 10 ⁻¹⁵
18	O ₃ + OH• → HO ₂ • + O ₂	2	1.6 × 10 ⁻¹²	940	6.8 × 10 ⁻¹⁴
19	OH• + OH• → O(³ P) + H ₂ O	2	4.2 × 10 ⁻¹²	240	1.9 × 10 ⁻¹²
20	OH• + OH• + M → H ₂ O ₂ + M	3			5.9 × 10 ⁻¹²
21	OH• + HO ₂ • → O ₂ + H ₂ O	2	4.8 × 10 ⁻¹¹	-250	1.1 × 10 ⁻¹⁰
22	H ₂ O ₂ + OH• → HO ₂ • + H ₂ O	2	2.9 × 10 ⁻¹²	160	1.7 × 10 ⁻¹²
23	O ₃ + HO ₂ • → OH• + O ₂ + O ₂	2	1.1 × 10 ⁻¹⁴	500	2.0 × 10 ⁻¹⁵
24	HO ₂ • + HO ₂ • → H ₂ O ₂ + O ₂	2	2.3 × 10 ⁻¹³	-600	1.7 × 10 ⁻¹²
25	HO ₂ • + HO ₂ • + M → H ₂ O ₂ + O ₂ + M ²	3	1.7 × 10 ⁻³³ [M]	-1000	4.9 × 10 ⁻³² [M]
26	O(³ P) + ClO• → Cl• + O ₂	2	3.0 × 10 ⁻¹¹	-70	3.8 × 10 ⁻¹¹
27	O(³ P) + OClO• → ClO• + O ₂	2	2.4 × 10 ⁻¹²	960	1.0 × 10 ⁻¹³
28	O(³ P) + HCl → Cl• + OH•	2	1.0 × 10 ⁻¹¹	3300	1.5 × 10 ⁻¹⁶
29	O(³ P) + HOCl → ClO• + OH•	2	1.7 × 10 ⁻¹³	0	1.7 × 10 ⁻¹³
30	OH• + Cl ₂ → Cl• + HOCl	2	1.4 × 10 ⁻¹²	900	6.7 × 10 ⁻¹⁴
31	OH• + ClO• → HCl + O ₂	2	1.5 × 10 ⁻¹²	-120	2.2 × 10 ⁻¹²
32	OH• + ClO• → HO ₂ • + Cl•	2	9.5 × 10 ⁻¹²	-120	1.4 × 10 ⁻¹¹
33	OH• + OClO• → HOCl + O ₂	2	4.5 × 10 ⁻¹³	-800	6.8 × 10 ⁻¹²
34	OH• + HCl → Cl• + H ₂ O	2	2.6 × 10 ⁻¹²	350	8.0 × 10 ⁻¹³
35	OH• + HOCl → ClO• + H ₂ O	2	3.0 × 10 ⁻¹²	500	5.0 × 10 ⁻¹³
36	HO ₂ • + Cl• → HCl + O ₂	2	1.8 × 10 ⁻¹¹	-170	3.2 × 10 ⁻¹¹
37	HO ₂ • + Cl• → OH• + ClO•	2	4.1 × 10 ⁻¹¹	450	9.1 × 10 ⁻¹²
38	HO ₂ • + ClO• → HOCl + O ₂	2	4.8 × 10 ⁻¹³	-700	5.0 × 10 ⁻¹²
39	O ₃ + Cl• → ClO• + O ₂	2	2.9 × 10 ⁻¹¹	260	1.2 × 10 ⁻¹¹
40	H ₂ O ₂ + Cl• → HO ₂ • + HCl	2	1.1 × 10 ⁻¹¹	980	4.1 × 10 ⁻¹³
41	Cl• + OClO• → ClO• + ClO•	2	3.4 × 10 ⁻¹¹	-160	5.8 × 10 ⁻¹¹
42	Cl• + HOCl → OH• + Cl ₂	2	2.5 × 10 ⁻¹²	130	1.6 × 10 ⁻¹²
43	ClO• + ClO• → Cl ₂ + O ₂	2	1.0 × 10 ⁻¹²	1590	4.8 × 10 ⁻¹⁵
44	ClO• + ClO• → ClOO• + Cl•	2	3.0 × 10 ⁻¹¹	2450	8.0 × 10 ⁻¹⁵
45	ClO• + ClO• → Cl• + Cl• + O ₂	2	3.5 × 10 ⁻¹³	1370	3.5 × 10 ⁻¹⁵
46	ClO• + ClO• + M → Cl ₂ O ₂ + M	3	See notes		3.4 × 10 ⁻¹³
47	Cl• + O ₂ + M → ClOO• + M	3	See notes		6.7 × 10 ⁻¹⁴
48	Cl• + ClOO• → Cl ₂ + O ₂	2	2.3 × 10 ⁻¹⁰	0	2.3 × 10 ⁻¹⁰
49	Cl• + ClOO• → ClO• + ClO•	2	1.2 × 10 ⁻¹¹	0	1.2 × 10 ⁻¹¹
50	Cl ₂ O ₂ + M → ClO• + ClO• + M	2	See notes		47
51	ClOO• + M → Cl• + O ₂ + M ³	2	See notes		2.7 × 10 ⁷
52	ClO• + O ₃ → ClOO• + O ₃	2			<1.4 × 10 ⁻¹⁷
53	ClO• + O ₃ → OClO• + O ₃	2			<1.0 × 10 ⁻¹⁸
54	Cl ₂ O + H ₂ O → HOCl + HOCl ³	2			<4.5 × 10 ⁻²³
55	HOCl + HOCl → Cl ₂ O + H ₂ O ³	2			<5.5 × 10 ⁻²²
56	Cl ₂ O + Cl• → Cl ₂ + ClO•	2	6.2 × 10 ⁻¹¹	-130	9.6 × 10 ⁻¹¹
57	Cl ₂ O + O(³ P) → ClO• + ClO•	2	2.7 × 10 ⁻¹¹	530	4.6 × 10 ⁻¹²
58	Cl ₂ O + ClO• → Cl ₂ + ClOO•	2			4.3 × 10 ⁻¹⁶
59	Cl ₂ O + ClO• → Cl ₂ + Cl• + O ₂	2			2.7 × 10 ⁻¹⁵
60	Cl ₂ + Wall → Loss	1			1.0 × 10 ⁻⁴
61	HOCl + Wall → Loss	1			1.0 × 10 ⁻³

Adapted from the works of *DeMore et al.* [1997] and *Sander et al.* [2000], except for (54)–(55) [Knauth et al., 1979], (58)–(59) [Basco and Dogra, 1971], and (60)–(61) measured directly in the chamber. k_{54} was measured at 323 K. $k_{55} = k_{54}/K_{\text{Cl}_2\text{O} + \text{H}_2\text{O}}$ at 298 K. $K_{\text{Cl}_2\text{O} + \text{H}_2\text{O}}$ at 298 K = 0.082 [Knauth et al., 1979]. [M] represents the number of molecules cm⁻³ of air. The rate constants for bimolecular reactions are given by $k = A \exp(-E/RT)$, where A is the Arrhenius factor in cm³ molecule⁻¹ s⁻¹ and E/R is the “activation temperature” in K. The rate constants for termolecular reactions (except for (25)) and thermal decomposition are calculated using the methods described by *DeMore et al.* [1997]. (1) The tabulated rate constants for termolecular reactions are the effective second-order rate constants assuming air at 1 atm as the third body. The photolysis and thermal decomposition rates (reactions (1)–(8) and (50)–(51)) are given in s⁻¹. (2) The overall rate constant for the reaction of self reaction of HO₂• is given by the sum of rates (24) and (25) multiplied by the factor: $1 + 1.4 \times 10^{-21} [\text{H}_2\text{O}] \exp(2200/T)$. (3) The rate constants for (52)–(55) are taken to match the upper limits.

phase and aqueous phase is determined in effect by the deviation from Henry’s law at the surface with respect to the bulk gas-phase concentration. Given that the concentration of an electrolytic species and its dissociation products are tracked by the model’s algorithm as a single entity, the overall (effective) Henry’s law constant is used

in the rate equations for mass transfer for the acid/base groups.

[21] In concentrated solutions, the activity of neutral species must also be considered when evaluating Henry’s law constants. Calculation of a Henry’s law constant for nonelectrolytes transported between the gas phase and

Table 4. Physical Solubility (Henry's Law) Constants and Accommodation Coefficients

Species	H_{298} M atm ⁻¹	$-\Delta H/R$ K	Reference	α	Reference
O ₃	0.011	2300	Kozak-Channing and Heltz [1983]	0.002	DeMore et al. [1997]
O ₂	1.3×10^{-3}	1500	Wilhelm et al. [1977]	0.01	Sander and Crutzen [1996] ^(E)
OH [•]	30	4500	Hanson et al. [1992]	0.1	Takami et al. [1998]
HO ₂ [•]	2.0×10^3		Schwartz [1984]	0.1	Hanson et al. [1992]
H ₂ O ₂	1.0×10^5	6300	Lind and Kok [1994]	0.1	DeMore et al. [1997]
HCl	1.1	2000	Marsh and McElroy [1985]	0.064	DeMore et al. [1997]
HOCl	930		Blatchley et al. [1992]	0.056	Sander and Crutzen [1996] ^(E)
Cl ₂	0.091	2500	Wilhelm et al. [1977]	0.07	Hu et al. [1995]
Cl ₂ O	17	1800	Wilhelm et al. [1977]	0.1	Estimated
CO ₂	0.034	2400	Sillen and Martell [1964]	0.01	Estimated

The temperature dependence for the Henry's law equilibrium constants is given by $H(T) = H_{298} \exp[(1/298 - 1/T) \times \Delta H/R]$ where ΔH is the enthalpy of dissolution. ^(E): estimated value.

aqueous phase is equivalent to determining the extent of salting-out as described in other environmental disciplines by Sechenow coefficients. Thus, for a nonelectrolytic species, the effective Henry's law constant is simply,

$$H^* = \left\{ \frac{1}{\gamma_A} \right\} H. \quad (6)$$

For a monoprotic electrolyte, the overall solubility of the gas-phase species HA is given by

$$H^* = \left\{ \frac{1}{\gamma_{HA}} + \frac{K_a}{\gamma_A \gamma_{H^+}} \right\} H \quad (7)$$

where K_a is the association/dissociation equilibrium constant. A similar expression can be derived for diprotic acids that dissociate sequentially in solution. A complete overview on the application of Henry's law coefficients in multiphase models is offered by Sander [1999].

[22] The accommodation coefficient, α , represents the probability of a collision of a gas-phase species with a droplet leading to successful transfer of the species from the gas phase to the aqueous phase. In general, the accommodation coefficient cannot be measured explicitly in the laboratory; rather, the overall uptake coefficient, also denoted by γ , is measured. The accommodation coefficient is decoupled from the other resistances to the overall uptake by applying models based on electrical circuit analogy [Jayne et al., 1990; Hu et al., 1995; Davidovits et al., 1995; Hanson, 1997]. These uptake resistance models, as well as the treatment of mass transfer processes in MAGIC, have a similar basis as the work of Schwartz [1986]. The reader is referred elsewhere for further discussion on accommodation and uptake coefficients [DeMore et al., 1997; Finlayson-Pitts and Pitts, 2000]. Table 4 provides values of the physical Henry's law constants and accommodation coefficients used in the present study.

2.4 Aqueous-Phase Chemistry

[23] The rates of aqueous-phase chemical reactions are selected from a compilation of the most recent kinetic data available. An initial mechanism is developed from the NIST database of solution-phase chemistry [Ross et al., 1998] and aqueous-phase inorganic radical chemistry reviews [Buxton et al., 1988; Neta et al., 1988]. Significant hydrogen peroxide concentrations are expected in the system, thus aqueous H₂O₂ photochemistry and thermal/H₂O₂-initiated

ozone decomposition mechanisms, including carbonate ion inhibition effects, were evaluated [Gonzalez and Martire, 1997; Sehested et al., 1991, 1998; Nemes et al., 2000]. Recently published aqueous chlorine free radical mechanisms are also considered [Buxton et al., 1998, 2000; Alegre et al., 2000], but aside from Cl₂ hydrolysis, nonradical chlorine chemistry is generally absent from available mechanisms. Finally, unfavorable kinetics and equilibria eliminate the necessity of including higher oxidation chlorine solutes, i.e., ClO_m[•] and ClO_{m-1}[•] where $m > 1$.

[24] Table 5 provides the rate constants for aqueous-phase chemical reactions relevant to the experimental conditions. Temperature dependences of many aqueous-phase reactions are not well established. Consequently, data are only provided for rate constants at 298 K. This does not pose any significant source of error since all experiments were performed near this temperature.

[25] Calculations of the light intensity in droplets based on Mie theory provide evidence that the actinic flux in aqueous particles may be enhanced, with respect to intensity, on average by a factor of two compared to the surrounding gas [Madronich, 1987; Ruggaber et al., 1997]. Such an enhancement would affect the photolysis rates inside the droplets. However, this effect is deemed of minimal importance in this study as the photolysis of O₃ occurs predominantly in the gas phase due to its low solubility. As confirmation, the photolysis rate of ozone in the presence of nonreactive sodium sulfate (Na₂SO₄) droplets measured in the experimental chamber was indeed equal to the photolysis rate in humid air alone. Therefore, the photolysis rates of aqueous-phase species are assumed equal to those of the corresponding gas-phase species.

[26] Acidic species included in the model possess electrolytic properties when transported into an aqueous medium. Other species, such as CO₂, can hydrolyze to form products that also dissociate in solution. For aqueous-phase reactions of atmospheric and laboratory systems, the timescale for ionization equilibrium is less than the timescale for chemical reaction [Schwartz and Freiberg, 1981]. Thus, it is assumed that electrolytic species and their dissociation products diffuse and react while preserving equilibrium and overall electroneutrality. Table 6 provides values for the equilibrium expressions used in this study.

2.5 Kinetic Salt Effect

[27] Although atmospheric models consider medium or salt effects in equilibrium/solubility expressions [Sander

Table 5. Aqueous-Phase Chemical Mechanism

No.	Reaction	O ^a	k ₂₉₈ ^b	Reference(s)
1001	H ₂ O ₂ + hν → 2 OH•	1	Photo	
1002	O ₃ + hν + H ₂ O → H ₂ O ₂ + O ₂	1	Photo	
1003	O ₃ + hν → O(³ P) + O ₂	1	Photo	
1004	O ₃ → O(³ P) + O ₂	1	3.0 × 10 ⁻⁶	Sehested et al. [1992]
1005	OH• + OH• → H ₂ O ₂	2	5.5 × 10 ⁹	Buxton et al. [1988, 2000]
1006	OH• + HO ₂ • → O ₂ + H ₂ O	2	1.0 × 10 ¹⁰	Elliot and Buxton [1992]
1007	OH• + O ₂ ⁻ → OH ⁻ + O ₂	2	1.0 × 10 ¹⁰	Elliot and Buxton [1992]
1008	OH• + H ₂ O ₂ → HO ₂ • + H ₂ O	2	2.7 × 10 ⁷	Buxton et al. [1988]
1009	OH• + HO ₂ ⁻ → HO ₂ • + OH ⁻	2	7.5 × 10 ⁹	Buxton et al. [1988]
1010	OH• + O ₃ → HO ₂ • + O ₂	2	1.1 × 10 ⁸	Neta et al. [1988]
1011	OH• + O ₃ ⁻ → O ₂ ⁻ + HO ₂ •	2	5.9 × 10 ⁹	Buxton et al. [1988]
1012	OH• + O ₃ ⁻ → O ₃ + OH ⁻	2	2.6 × 10 ⁹	Buxton et al. [1988]
1013	O• ⁻ + O ₂ → O ₃ ⁻	2	3.6 × 10 ⁹	Buxton et al. [1988]
1014	O• ⁻ + HO ₂ ⁻ → O ₂ ⁻ + OH ⁻	2	4.0 × 10 ⁸	Buxton et al. [1988]
1015	O• ⁻ + O ₂ ⁻ + H ₂ O → 2 OH ⁻ + O ₂	2	6.0 × 10 ⁸	Buxton et al. [1988]
1016	O• ⁻ + OH• → HO ₂ ⁻	2	≤ 2.0 × 10 ¹⁰	Buxton et al. [1988]
1017	O• ⁻ + O• ⁻ + H ₂ O → HO ₂ ⁻ + OH ⁻	2	1.0 × 10 ⁹	Buxton et al. [1988]
1018	O• ⁻ + O ₃ → O ₂ ⁻ + O ₂	2	1.0 × 10 ⁸	Gonzalez and Martire [1997]
1019	O• ⁻ + O ₃ ⁻ → 2 O ₂ ⁻	2	7.0 × 10 ⁸	Buxton et al. [1988]
1020	HO ₂ • + HO ₂ • → H ₂ O ₂ + O ₂	2	8.3 × 10 ⁵	Bielski [1978]
1021	HO ₂ • + O ₂ ⁻ → HO ₂ ⁻ + O ₂	2	1.0 × 10 ⁸	Bielski [1978]
1022	HO ₂ • + H ₂ O ₂ → OH• + O ₂ + H ₂ O	2	0.5	Weinstein and Bielski [1979]
1023	HO ₂ • + O ₃ → OH• + 2 O ₂	2	< 1.0 × 10 ⁴	Sehested et al. [1984b]
1024	O ₂ ⁻ + O ₂ ⁻ + 2 H ₂ O → H ₂ O ₂ + 2 OH ⁻ + O ₂	2	< 0.3	Bielski [1978]
1025	O ₂ ⁻ + H ₂ O ₂ → OH• + OH ⁻ + O ₂	2	0.13	Weinstein and Bielski [1979]
1026	O ₂ ⁻ + O ₃ → O ₃ ⁻ + O ₂	2	1.5 × 10 ⁹	Sehested et al. [1983]
1027	O ₂ ⁻ + O ₃ ⁻ → 2 OH ⁻ + 2 O ₂	2	5.0 × 10 ⁴	Gonzalez and Martire [1997]
1028	O ₃ ⁻ → O• ⁻ + O ₂	1	5.0 × 10 ³	Neta et al. [1988]
1029	O ₃ ⁻ + H ⁺ → OH• + O ₂	2	9.0 × 10 ⁹	Sehested et al. [1984a]
1030	O ₃ ⁻ + H ⁺ → HO ₃ •	2	5.0 × 10 ¹⁰	Neta et al. [1988]
1031	HO ₃ • → OH• + O ₂	1	1.0 × 10 ⁵	Neta et al. [1988]
1032	HO ₃ • → O ₃ ⁻ + H ⁺	1	3.7 × 10 ⁴	Bühler et al. [1984]
1033	O ₃ + OH ⁻ → HO ₂ • + O ₂	2	48	Neta et al. [1988]
1034	O ₃ + OH ⁻ → HO ₂ • + O ₂ ⁻	2	70	Neta et al. [1988]
1035	O ₃ + H ₂ O ₂ → OH• + HO ₂ • + O ₂	2	0.025	Sehested et al. [1992]
1036	O ₃ + HO ₂ ⁻ → HO ₃ • + O ₂ ⁻	2	5.5 × 10 ⁶	Neta et al. [1988]
1037	O(³ P) + O ₂ → O ₃	2	4.0 × 10 ⁹	Klaening et al. [1984]
1038	O(³ P) + OH ⁻ → HO ₂ ⁻	2	4.2 × 10 ⁸	Sauer et al. [1984]
1039	O(³ P) + H ₂ O ₂ → OH• + HO ₂ •	2	1.6 × 10 ⁹	Sauer et al. [1984]
1040	O(³ P) + HO ₂ ⁻ → OH• + O ₂ ⁻	2	5.3 × 10 ⁹	Sauer et al. [1984]
1041	O• ⁻ + H ₂ O ₂ → O ₂ ⁻ + H ₂ O	2	≤ 5.0 × 10 ⁸	Buxton et al. [1988]
1101	HOCl + hν → OH• + Cl•	1	Photo	
1102	ClO ⁻ + hν → Cl• + O(³ P)	1	Photo	
1103	ClO ⁻ + hν → Cl• + O• ⁻	1	Photo	
1104	ClO ⁻ + hν + H ₂ O → Cl• + H ₂ O ₂	1	Photo	
1105	Cl ₂ + hν → 2 Cl•	1	Photo	
1106	Cl• + Cl ⁻ → Cl ₂ ⁻	2	8.5 × 10 ⁹	Buxton et al. [1998]
1107	Cl ₂ ⁻ → Cl• + Cl ⁻	1	6.0 × 10 ⁴	Buxton et al. [1998]
1108	OH• + Cl ⁻ → HOCl• ⁻	2	4.3 × 10 ⁹	Jayson et al. [1973]
1109	HOCl• ⁻ → OH• + Cl ⁻	1	6.1 × 10 ⁹	Jayson et al. [1973]
1110	Cl ₂ ⁻ + H ₂ O → HOClH• + Cl ⁻	1	1.3 × 10 ³	Buxton et al. [1998]
1111	HOClH• + Cl ⁻ → Cl ₂ ⁻ + H ₂ O	2	8.0 × 10 ⁹	Alegre et al. [2000]
1112	Cl• + H ₂ O → HOClH•	1	2.5 × 10 ⁵	Buxton et al. [1998]
1113	HOClH• → Cl• + H ₂ O	1	5.0 × 10 ⁴	Alegre et al. [2000]
1114	HOCl• ⁻ + H ⁺ → HOClH•	2	3.0 × 10 ¹⁰	Alegre et al. [2000]; Jayson et al. [1973]
1115	HOClH• → HOCl• ⁻ + H ⁺	1	1.0 × 10 ⁸	Alegre et al. [2000]; McElroy [1990]
1116	Cl ₂ ⁻ + Cl ₂ ⁻ → Cl ₂ + 2 Cl ⁻	2	1.8 × 10 ⁹	Jacobi et al. [1999]
1117	Cl ₂ ⁻ + OH ⁻ → HOCl• ⁻ + Cl ⁻	2	4.0 × 10 ⁶	Jacobi et al. [1997]
1118	HOCl• ⁻ + Cl ⁻ → Cl ₂ ⁻ + OH ⁻	2	1.0 × 10 ⁴	Grigor'ev et al. [1987]
1119	Cl ₂ ⁻ + HO ₂ • → 2 Cl ⁻ + H ⁺ + O ₂	2	1.0 × 10 ⁹	Navaratnam et al. [1980]
1120	Cl ₂ ⁻ + O ₂ ⁻ → 2 Cl ⁻ + O ₂	2	1.0 × 10 ⁹	Neta et al. [1988]
1121	Cl ₂ ⁻ + H ₂ O ₂ → 2 Cl ⁻ + HO ₂ • + H ⁺	2	5.0 × 10 ⁴	Elliot [1989]
1122	Cl ₂ ⁻ + OH• → HOCl + Cl ⁻	2	1.0 × 10 ⁹	Wagner et al. [1986]; Buxton et al. [2000]
1123	Cl ₂ ⁻ + O ₃ → ClO ⁻ + Cl + O ₂	2	9.0 × 10 ⁷	Bielski [1993]
1124	Cl• + Cl• → Cl ₂	2	8.8 × 10 ⁷	Neta et al. [1988]
1125	Cl• + OH ⁻ → HOCl• ⁻	2	1.8 × 10 ¹⁰	Neta et al. [1988]
1126	Cl• + HO ₂ • → Cl ⁻ + H ⁺ + O ₂	2	3.1 × 10 ⁹	Graedel and Goldberg [1983]
1127	Cl• + H ₂ O ₂ → Cl ⁻ + HO ₂ • + H ⁺	2	4.1 × 10 ⁷	Graedel and Goldberg [1983]
1128	Cl ₂ + HO ₂ • → Cl ₂ ⁻ + H ⁺ + O ₂	2	1.0 × 10 ⁹	Bjergbakke et al. [1981]
1129	HOCl + O ₂ ⁻ → Cl ⁻ + OH• + O ₂	2	7.5 × 10 ⁶	Long and Bielski [1980]
1130	Cl ⁻ + O ₃ → ClO ⁻ + O ₂	2	2.0 × 10 ⁻³	Neta et al. [1988]
1131	ClO ⁻ + O ₃ → Cl ⁻ + 2 O ₂	2	110	Neta et al. [1988]
1132	Cl ₂ + H ₂ O → HOCl + Cl ⁻ + H ⁺	1	22	Wang and Margerum [1994]

Table 5. (continued)

No.	Reaction	O ^a	k ₂₉₈ ^b	Reference(s)
1133	HOCl + Cl ⁻ + H ⁺ → Cl ₂ + H ₂ O	3	2.1 × 10 ⁴	Wang and Margerum [1994]
1134	Cl ₂ + OH ⁻ → HOCl + Cl ⁻	2	1.0 × 10 ⁸	Estimate, this work (see sections 4.1 and 4.2)
1135	HOCl + Cl ⁻ → Cl ₂ + OH ⁻	2	2.0 × 10 ⁻³	Estimate, this work (using k ₁₁₃₄ and equilibrium constants)
1136	Cl ₂ + H ₂ O ₂ → 2 Cl ⁻ + 2 H ⁺ + O ₂	2	See notes	Held et al. [1978]; Connick [1947]; Makower and Bray [1933]
1137	Cl ₂ + HO ₂ ⁻ → 2 Cl ⁻ + H ⁺ + O ₂	2	1.1 × 10 ⁸	Davis et al. [1992]; Ruiz-Ibanez and Sandall [1991]
1138	HOCl + HO ₂ ⁻ → Cl ⁻ + O ₂ + H ₂ O	2	4.4 × 10 ⁷	Held et al. [1978]
1139	Cl ⁻ + H ₂ O ₂ → ClO ⁻ + H ₂ O	2	1.8 × 10 ⁻⁹	Hansen and Espenson [1995]; Mohammad and Liebhafsky [1934]
1140	Cl ⁻ + H ₂ O ₂ + H ⁺ → HOCl + H ₂ O	3	8.3 × 10 ⁻⁷	Hansen and Espenson [1995]; Mohammad and Liebhafsky [1934]
1141	Cl ₂ O + H ₂ O → 2 HOCl	1	7.0	Beach and Margerum [1990]
1142	HOCl + HOCl → Cl ₂ O + H ₂ O	2	0.08	Beach and Margerum [1990]
1143	Cl ₂ O + H ⁺ + H ₂ O → HOCl + HOCl + H ⁺	2	270	Beach and Margerum [1990]; Swain and Crist [1972]
1144	HOCl + HOCl + H ⁺ → Cl ₂ O + H ⁺ + H ₂ O	3	3.1	Swain and Crist [1972]
1145	Cl ₂ O + hν → 2 HCl + O ₂	1	Photo	Johnsson et al. [1993]
1201	HCO ₃ ⁻ + OH [•] → CO ₃ ^{•-} + H ₂ O	2	8.5 × 10 ⁶	Buxton et al. [1988]
1202	CO ₃ ^{•-} + OH [•] → CO ₃ ^{•-} + OH ⁻	2	3.9 × 10 ⁸	Buxton et al. [1988]
1203	HCO ₃ ⁻ + O ₂ ^{•-} → CO ₃ ^{•-} + HO ₂ ⁻	2	1.5 × 10 ⁶	Schmidt [1972]
1204	CO ₃ ^{•-} + O ₂ ^{•-} → CO ₃ ^{•-} + 2 OH ⁻	2	<5.0 × 10 ⁵	Gonzalez and Martire [1997]
1205	CO ₃ ^{•-} + H ₂ O ₂ → HCO ₃ ⁻ + HO ₂ [•]	2	4.3 × 10 ⁵	Draganić et al. [1991]
1206	CO ₃ ^{•-} + HO ₂ ⁻ → CO ₃ ^{•-} + HO ₂ [•]	2	3.0 × 10 ⁷	Gonzalez and Martire [1997]
1207	CO ₃ ^{•-} + HO ₂ [•] → HCO ₃ ⁻ + O ₂	2	5.6 × 10 ⁷	Behar et al. [1970]
1208	CO ₃ ^{•-} + O ₂ ^{•-} → CO ₃ ^{•-} + O ₂	2	6.5 × 10 ⁸	Ericksen et al. [1985]
1209	CO ₃ ^{•-} + O ₃ ^{•-} → CO ₃ ^{•-} + O ₃	2	6.0 × 10 ⁷	Holcman et al. [1982]
1210	CO ₃ ^{•-} + CO ₃ ^{•-} + O ₂ + 2 H ₂ O → 2 CO ₂ •H ₂ O + 2 O ₂ ^{•-}	2	2.2 × 10 ⁶	Huie and Clifton [1990]
1211	HCO ₃ ⁻ + Cl [•] → CO ₃ ^{•-} + Cl ⁻ + H ⁺	2	2.2 × 10 ⁸	Mertens and von Sonntag [1995]
1212	CO ₃ ^{•-} + Cl [•] → CO ₃ ^{•-} + Cl ⁻	2	5.0 × 10 ⁸	Mertens and von Sonntag [1995]
1213	CO ₃ ^{•-} + Cl ₂ ^{•-} → CO ₃ ^{•-} + 2 Cl ⁻	2	2.7 × 10 ⁶	Estimate, Herrmann et al. [2000]
1214	Cl ₂ + HCO ₃ ⁻ + H ₂ O → HOCl + Cl ⁻ + CO ₂ •H ₂ O	2	3.2 × 10 ³	Estimate, this work (see sections 4.1 and 4.2)
1215	Cl ₂ + CO ₃ ^{•-} + H ₂ O → HOCl + Cl ⁻ + HCO ₃ ⁻	2	1.4 × 10 ⁵	Estimate, this work (see sections 4.1 and 4.2)
1216	HOCl + Cl ⁻ + CO ₂ •H ₂ O → Cl ₂ + HCO ₃ ⁻ + H ₂ O	3	2.7	Estimate, this work (using k ₁₂₁₄ and equilibrium constants)
1217	HOCl + Cl ⁻ + HCO ₃ ⁻ → Cl ₂ + CO ₃ ^{•-} + H ₂ O	3	0.012	Estimate, this work (using k ₁₂₁₅ and equilibrium constants)
1218	CO ₂ •H ₂ O + OH [•] → CO ₃ ^{•-} + H ⁺ + H ₂ O	2	≥ 7.0 × 10 ⁴	Czapski et al. [1999]

$$k_{1136} = 183.3/(2.27[\text{H}^+][\text{Cl}^-] + 1).$$

^aReaction order.

^bUnits in M^(1 - O) s⁻¹.

and Crutzen, 1996; Moldanova and Ljungstrom, 2001], kinetic salt effects are also warranted in the calculation of the rate constants of aqueous-phase reactions. The historical and theoretical background of the kinetic salt effect is found in classic review articles [Rochester, 1971; Perlmutter-Hayman, 1971; Pethybridge and Prue, 1972]. Briefly, the kinetics of a reaction between two (or more) solution-phase reactants is dictated by the formation of an activated complex, X, in accordance with transition state theory, as shown in equation (8) for the case of a bimolecular reaction between A and B.



[28] Bronsted and Bjerrum independently proposed that the rate of aqueous-phase chemical reactions depends not

only on the activities of the reactants, but also on the activity of a “critical complex,” i.e., the activated complex. The rate constant for an aqueous-phase chemical reaction is determined by the so-called Bronsted–Bjerrum relation:

$$k = k^\circ F = k^\circ \frac{\gamma_A \gamma_B}{\gamma_X} \quad (9)$$

where k° is the rate constant at infinite dilution while γ_A , γ_B , and γ_X denote the activity coefficients of the reactants and the activated complex, respectively. The ratio of the activity coefficients, F , is the kinetic activity factor.

[29] From simple application of the Debye limiting law, the Bronsted–Bjerrum relation becomes

$$\log k = \log k^\circ + A \Delta z^2 \sqrt{I} \quad (10)$$

Table 6. Aqueous-Phase Acid/Base Equilibrium Expressions

Equilibrium reaction	K_{a298}	$-\Delta H/R$	Reference
	M	K	
H ₂ O ⇌; H ⁺ + OH ⁻	1.00 × 10 ⁻¹⁴	6710	Smith and Martell [1976]
H ₂ O ₂ ⇌; H ⁺ + HO ₂ ⁻	2.50 × 10 ⁻¹²		Sauer et al. [1984]
HO ₂ [•] ⇌; H ⁺ + O ₂ ^{•-}	2.05 × 10 ⁻⁵		Bielski [1978]; Bielski et al. [1985]
OH [•] ⇌; H ⁺ + O ^{•-}	1.26 × 10 ⁻¹²		Buxton et al. [1988]
HCl ⇌; H ⁺ + Cl ⁻	1.70 × 10 ⁶	6896	Marsh and McElroy [1985]
HOCl ⇌; H ⁺ + ClO ⁻	2.80 × 10 ⁻⁸		Huthwelker et al. [1995]
CO ₂ •H ₂ O ⇌; H ⁺ + HCO ₃ ⁻	4.30 × 10 ⁻⁷	920	Sillén and Martell [1964]
HCO ₃ ⁻ ⇌; H ⁺ + CO ₃ ^{•-}	4.30 × 10 ⁻¹¹	1786	Sillén and Martell [1964]
HCO ₃ [•] ⇌; H ⁺ + CO ₃ ^{•-}	≥ 1		Czapski et al. [1999]

The temperature dependence for the equilibrium constants is given by $K_a(T) = K_{a298} \exp[(1/298 - 1/T) \times \Delta H/R]$ where ΔH is the enthalpy of reaction.

where

$$\Delta z^2 = z_X^2 - z_A^2 - z_B^2 - \dots \quad (11)$$

In the above equations, z represents the charge of the species, A is the Debye parameter, and I is the ionic strength of the solution. Equation (10), commonly known as the Debye–Bronsted equation, simplifies in the case of a bimolecular reaction to

$$\log k = \log k^\circ + 2Az_A z_B \sqrt{I} \quad (12)$$

[30] The validity of equation (12) at ionic strengths below 0.1 M is manifest by the Livingston diagram [Benson, 1960; Finlayson-Pitts and Pitts, 2000]. The range of ionic strength is enhanced using the extended Debye–Huckel limiting law, yielding the Debye–Huckel–Bronsted equation:

$$\log k = \log k^\circ + \frac{A\Delta z^2 \sqrt{I}}{1 + \sqrt{I}} \quad (13)$$

[31] A revised Livingston diagram shows that equation (13) yields good agreement for $I \leq 0.5$ M [Pethybridge and Prue, 1972]. However, at relative humidities near the deliquescence point, marine aerosol can be expected to reach ionic strengths near 6 M. Some progress, at least in experimental settings, can be made by further extending the Debye–Huckel limiting law as proposed by Guggenheim and incorporating it into the Bronsted–Bjerrum relation yielding

$$\log k = \log k^\circ + \frac{A\Delta z^2 \sqrt{I}}{1 + \sqrt{I}} + \sum_i B_{A,i} C_i + \sum_j B_{B,j} C_j - \sum_k B_{X,k} C_k \quad (14)$$

where $B_{m,n}$ are specific interaction coefficients between ion m and all ions of opposite charge n having concentration C_n . This expression can be simplified if the relative ionic concentrations remain fairly unchanged in a series of experiments into

$$\log k = \log k^\circ + \frac{A\Delta z^2 \sqrt{I}}{1 + \sqrt{I}} + B \quad (15)$$

where the empirical kinetic salt effect parameter B is separated into the individual contributions B_A , B_B and B_X due to the reactants and the activated complex:

$$B = B_A + B_B - B_X \quad (16)$$

Equation (15) is also known as the Debye–Huckel–Bronsted–Davies equation when B is substituted by an equivalent parameter b' that is written at times with an opposite sign:

$$\log k = \log k^\circ + \frac{A\Delta z^2 \sqrt{I}}{1 + \sqrt{I}} - b' I \quad (17)$$

The range of B (or $-b'$) varies greatly with reported values reaching as high and positive as 0.74 [Kumar *et al.*, 1986]

and as large and negative as -0.86 [Pethybridge and Prue, 1972]. An important corollary of the theory above is that for neutral–neutral or neutral–ion reactions equation (15) reduces to

$$\log k = \log k^\circ + BI \quad (18)$$

implying that, at sufficiently high ionic strengths, a kinetic salt effect may also be present.

[32] Unfortunately, the determination of the kinetic salt effect in atmospheric modeling applications is cumbersome due primarily to the following physical processes: (1) Changes in relative humidity alter the overall concentration and ionic strength of an aqueous aerosol particle. (2) Absorption and volatilization of constituents from and to the gas phase affect the relative concentrations of ionic species with respect to each other. (3) Chemical reactions within the droplet (or at the surface of the droplet) also affect the overall composition of the aerosol.

[33] At best, atmospheric models can only attempt to estimate the kinetic salt effect. This can be done by several methods. First, the kinetic activity factor, F , may be evaluated directly through calculation of individual activity coefficients. However, the activities of many organic species, inorganic constituents at atypical oxidation states, and radical ions cannot be calculated with current activity coefficient models thus leading to their estimation by extended Debye–Huckel expressions [Sander and Crutzen, 1996; Moldanova and Ljungstrom, 2001]. In addition, the activity of the activated complex must also be estimated.

[34] Second, the effective rate constant may be determined via Debye–Huckel–Bronsted equations. The advantage, simplicity, is easily overshadowed by the disadvantage, lack of validity beyond ionic strengths of 0.5 M. Nonetheless, if insight of the importance of various processes, not absolute values, is the desired output from the modeling studies, such a method may prove sufficient and attractive. With the understanding of this important caveat, this is the method employed in the current implementation of MAGIC.

[35] Finally, the effective rate constant may be determined via Debye–Huckel–Bronsted–Guggenheim—or the similar Debye–Huckel–Bronsted–Davies—equations. Considering equation (14) in an experimental context, the determination of the specific interaction parameters $B_{m,n}$ —primarily for the activated complexes and radical ions—could be an unapproachably difficult task. Instead, determination of an overall kinetic salt effect parameter, B , is more advantageous. Experimental studies shall determine whether this parameter is chosen as a constant coefficient to ionic strength, or adjusted as a function of the concentrations of the primary ions in solution.

[36] If a compilation of aqueous-phase chemical data for atmospheric modeling is commissioned [Jacob, 2000], the selection of a common method for determining the kinetic salt effect should be included and suggestions to quantify such effects conveyed accordingly.

2.6. Aqueous-Phase Diffusion

[37] Only species actively transported between the gas phase and aqueous phase, and their hydrolysis and/or dissociation products, need to be considered as presenting

spatial gradients within the droplets [Jacob, 1986]. For these species, a concentration value at the surface must be determined in order to calculate volatilization fluxes and another average bulk concentration value must be determined to calculate the overall rate of aqueous-phase reactions within the droplet. All other species are assumed well mixed throughout the droplet.

[38] The temporal and spatial dependence of the concentration of species undergoing diffusion and irreversible reaction inside a spherically symmetrical droplet is governed by the equations

$$\frac{dC_j(r, t)}{dt} = D_{aq, j} \frac{1}{r^2} \frac{d}{dr} \left(r^2 \frac{dC_j(r, t)}{dr} \right) - k_{L, j}(r, t) C_j(r, t) + P_j(r, t). \quad (19)$$

where r is the distance from the center of the droplet; t is the temporal variable; $D_{aq, j}$ and $C_j(r, t)$ are the aqueous-phase diffusivity and concentration of species j , respectively; and $R(r, t)$ is the aqueous-phase reaction rate. Equation (19) also contain two chemical reaction terms: an effective global first-order loss rate constant $k_{L, j}$ (s⁻¹) for species j and a global production rate P_j (M s⁻¹) by aqueous-phase reactions both of which may depend on other species that present radial dependencies on concentration, i.e., other species transported to and from the gas phase.

[39] In the extreme situation of a droplet suddenly exposed to a gas-phase reagent whose concentration remains constant at the surface (in the absence of any aqueous-phase production), the rate of uptake into the droplet calculated from the steady state solution does not greatly exceed the rate calculated from the time dependent solution [Schwartz and Freiberg, 1981]. Accordingly, through application of the steady state solution, equation (19) become

$$D_{aq, j} \frac{1}{r^2} \frac{d}{dr} \left(r^2 \frac{dC_j(r)}{dr} \right) - k_{L, j}(r) C_j(r) + P_j(r) = 0, \quad (20)$$

subject to the boundary conditions

$$\left(\frac{dC_j(r)}{dr} \right)_{r=0} = 0 \quad (21)$$

$$C_j(R_p) = C_j^*, \quad (22)$$

where C_j^* is the concentration at the droplet surface. After mathematical manipulation of equations (20)–(22), the average bulk concentrations of aqueous-phase species are determined by solving a system of closely coupled nonlinear equations given by:

$$\langle C_j \rangle = QC_j^* + (1 - Q) \frac{P_j}{k_{L, j}}. \quad (24)$$

where

$$Q = 3 \left(\frac{\coth q_j}{q_j} - \frac{1}{q_j^2} \right) \quad (25)$$

and

$$q_j = R_p \left(\frac{k_{L, j}}{D_{aq, j}} \right)^{1/2}. \quad (26)$$

2.7. Interfacial Reactions

[40] The rates of gas-phase reactions may be enhanced by the presence of liquid or solid surfaces [Ravishankara, 1997]. A thorough analysis of the data (see results section) shows that in the experimental system at hand, the prospect of a surface reaction involving chloride ions and hydroxyl radicals is a fully reasonable deduction. A brief description of the modeling of the surface mechanism follows.

[41] Results by Knipping *et al.* [2000] suggest that chloride ions present at aqueous surface hold a high affinity for gas-phase OH• radicals, thereby enhancing scavenging of the gas at the surface and leading to the formation of a surface complex species, OH•••Cl_{surface}⁻,



Recalling the equation for the molecular collision rate per surface area, the formation of OH•••Cl_{surface}⁻ is determined by

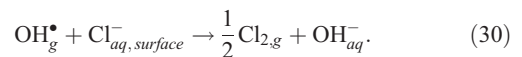
$$R_{27} = \gamma_s \frac{\bar{c}}{4} A [\text{OH}^\bullet]_g, \quad (28)$$

where γ_s is the overall probability of Reaction 27 occurring once an OH• molecule strikes the droplet surface and A is the total particle surface area. The probability, γ_s , is a function of the availability of chloride ions at the surface and the likelihood of complexation. Molecular dynamic simulations [Jungwirth and Tobias, 2000] suggest that the fraction of the droplet surface covered by chloride ions can be estimated as 0.02[Cl⁻], where [Cl⁻] is the concentration of chloride expressed in M (mol L⁻¹), and the numerical constant has units of M⁻¹. The probability is therefore be expressed as

$$\gamma_s = 0.02 \gamma'_s [\text{Cl}^-], \quad (29)$$

where γ'_s is a fitting parameter incorporating the probability of an OH• radical forming the complex. The value of the γ'_s term is explored in the results section.

[42] In contrast to earlier studies of the same system [Knipping *et al.*, 2000], instead of attempting to speculate the specific fate of OH•••Cl_{surface}⁻ and the kinetics of the individual reactions thereof, the present study is limited to finding an overall rate expression on the premise that the formation of the chloride–hydroxyl complex is the rate limiting step of an overall reaction given by



This assumption agrees with initial modeling studies and offers several important advantages: (1) an ability to focus solely on the leading hypothesis: a surface enhanced mechanism determines the rate of chlorine formation; (2) provides a straightforward expression amenable to inclusion by other models; and (3) underscores the possibility of

several, not necessarily exclusive, reactions in the interfacial region involving the initially formed $\text{OH}\bullet\bullet\bullet\text{Cl}_{\text{surface}}^{\bullet-}$.

2.8. Governing Equations

[43] Applying the results obtained from mass transfer theory of Schwartz [1986] in combination with terms for bulk gas-phase chemical reaction, R_g , and interfacial reactions, R_{int} , the temporal variations of the concentrations of gas-phase and aqueous-phase species are described by

$$\frac{dC_g}{dt} = -k_{mt}w_L C_g + \frac{k_{mt}}{HRT} C_{aq}w_L + R_g \pm R_{int} \quad (31)$$

$$\frac{dC_{aq}}{dt} = k_{mt}C_g - \frac{k_{mt}}{HRT} C_{aq} + \langle R_{aq} \rangle \mp \frac{R_{int}}{w_L}, \quad (32)$$

where R_{int} is the overall rate of the interfacial mechanism as expressed in equation (28).

2.9. Methodology of Mechanism Validation

[44] An extensive literature search performed in order to assemble the chemical mechanism for the “relatively simple” system of interest, resulted in nearly 200 parameters including aqueous-phase and gas-phase rate constants, Henry’s law constants, accommodation coefficients and acid/base equilibrium constants. Although many of these parameters are well known, the potential for error is explored through sensitivity studies.

[45] The strategy developed in collaboration with our experimental and theoretical chemistry associates in order to clarify the mechanism for Cl₂ production is as follows:

1. Create a model that attempts to reproduce the experimental results using the best known data and representation of the physical and chemical processes of the experiments.

If the model cannot duplicate, within reasonable error, the levels of chlorine detected in the experimental chamber the ensuing steps were taken:

2. Study the sensitivity of the system in order to determine which reactions are significant and evaluate whether the uncertainty of these reactions can explain the Cl₂ production.

3. Investigate whether error in model ino5g (initial conditions and aerosol or chamber parameters) can justify for the level of discrepancy.

4. Investigate whether any unknown, yet rational, chemical reaction or physical process solely in the aqueous phase or gas phase can account for the Cl₂ formation.

It is only upon complete failure of all three tests (2–4) above that the likelihood for chemical reaction(s) at the interface is (are) considered. The burgeoning evidence for these types of reactions suggests that a similar reaction is plausible in our system (see section 1). The methodology invoked in the mechanism’s analysis proceeds by performing the following tests:

5. Determine which of the reactive gas concentrations, in combination with the amount of surface area available, correlates better with the observed rates of Cl₂ formation.

6. Confirm the likelihood of the participation of the preferred gas-phase species using theoretical and thermodynamic considerations.

[46] If the above tests yield a potential candidate for a surface reaction, then a surface mechanism can be reason-

ably inferred. Given the level of complexity involved in determining such a mechanism from limited laboratory data, it is preferable to determine a kinetic expression for the overall rate of Cl₂ production. Nevertheless, in order to support such a hypothesis, at least one compatible and reasonable mechanism must be presented in a qualitative fashion. Two such mechanisms that are chemically sound given the present level of understanding of interfacial processes on aqueous sodium chloride surfaces are presented.

[47] Finally, once an overall rate expression is obtained, this “interface model” is subjected to the same rigorous tests as the “noninterface model.” In this case, we wish not only to see whether any reasonable discrepancy between the model results and experimental data can be explained by variations in model input and/or parameters, but also to guarantee that the level of modeled Cl₂ production is not attributable to some fortuitous combination of highly sensitive parameters.

[48] Given the vanguard nature of these combined theoretical, experimental and kinetic modeling efforts, the results presented here are intended to serve as a springboard to further investigate a promising area in aerosol reactions in the atmosphere, i.e., reactions at the air–water interface.

3. Experimental Modeling

3.1. Established Chemical Mechanisms

[49] In order to clarify the mechanism for the observed Cl₂ production in the chamber experiments, different pathways are evaluated by performing simulations using MAGIC and comparing the results to the experimental data. Table 7 contains the parametric data of three principal experimental scenarios. The experiments are denoted as the low, middle and high O₃ experiments according to their initial ozone concentration. The model, in correspondence with experimental procedure, runs for 12 min in the dark, allowing the system to approach an initial equilibrium state. Determination of Cl₂ levels in the chamber after each 2-min photolysis period (10 total) requires flowing some of the contents through the atmospheric pressure ionization mass spectrometer (API-MS) during a 6-min dark sampling period, effectively diluting the chamber. In order to reproduce the chamber conditions, this effect is modeled as a first-order loss resulting in a decrease in gas-phase concentrations in the proportion given by the dilution factor on Table 7. The first-order dilution rate constant is given by

$$k_{\text{dilution}} = -\frac{\log(\text{dilution factor})}{t_{\text{API Sampling}}} \quad (33)$$

[50] Modeled Cl₂ and O₃ time series (Figures 2–7) exhibit a sawtooth shape indicative of API gas-phase dilution; experimental Cl₂ measurements are shown as 6-min bars corresponding to the average concentration measured during sampling, whereas experimental O₃ measurements are shown as instantaneous points measured at the onset of the API dark sampling period.

[51] O₂, N₂ and H₂O concentrations do not change during the sampling period due to the addition of humid ultrapure air. The aerosol particle concentration, i.e., the mass of aerosol particles per unit volume in the chamber, also does not change during the simulation. The count and measurement of the aerosol population is performed twice before

Table 7. Model Parameters for Sequence of Experiments Varying Initial O₃ Concentration

Parameter	Units	Simulation		
		High ozone	Middle ozone	Low ozone
$[\text{O}_3]_0$	molecule cm^{-3}	3.37×10^{14}	2.44×10^{14}	6.31×10^{13}
$[\text{CO}_2]_0$	ppm	17.7	16.6	16.9
$[\text{NaCl}]_0$	M	4.30	4.30	4.28
Dilution factor		0.945	0.940	0.956
RH	%	81.8	81.8	81.9
Temperature	K	298	298	298
Number	cm^{-3}	1.86×10^5	1.43×10^5	1.09×10^5
Median diameter	nm	224	227	209
Mass transfer radius	nm	396	353	353
Geometric standard deviation	—	1.93	1.97	2.09
Area	nm^2/cm^3	8.16×10^{10}	5.79×10^{10}	4.42×10^{10}
Volume	nm^3/cm^3	9.74×10^{12}	6.92×10^{12}	5.95×10^{12}

The mass transfer radius is the effective radius used in the quasi-size-dependent calculation of the mass transfer rate coefficient, k_m , and is derived by integrating over the droplet population [Schwartz, 1986; Sander and Crutzen, 1996].

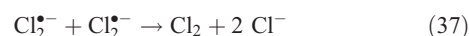
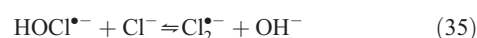
initiating the experiments and twice after completing the final API sampling. An average aerosol distribution is used throughout the model simulation; the error associated with this assumption is treated later in this discussion.

[52] Evaluation of the full mechanism based on known gas-phase and aqueous-phase chemistry is equivalent to an evaluation of the reaction pathway proposed by *Oum et al.* [1998] (see Figure 1). Minor differences reflect the current understanding regarding the equilibria formed by the oxidation of Cl^- by $\text{OH}^\bullet$ in aqueous solution and the reactions of $\text{Cl}^\bullet/\text{Cl}_2^\bullet-$ with water [Jayson *et al.*, 1973; Wagner *et al.*, 1986; McElroy, 1990; Buxton *et al.*, 1998; Alegre *et al.*, 2000]. The source of this mechanism begins with the photolysis of ozone and subsequent reaction of excited oxygen atoms, $\text{O}(^1\text{D})$, with water, leading to hydroxyl radical generation in the gas phase. $\text{OH}^\bullet$ is then taken up by the particles and oxidation of the chloride ion eventually leads to Cl_2 release into the gas phase. These simulations based on established chemistry and mass transfer parameters are designated as “base case” simulations.

[53] Figure 2 shows base case modeling results, corresponding to the middle ozone experiment, using only established gas-phase and aqueous-phase chemical reactions and mass transfer processes pertinent to the simulated system—no interfacial reactions have been modeled. Although the ozone decay is in agreement with experimental data (as expected since the photolysis rate was measured independently and input directly into the model), the amount of Cl_2 predicted is more than three orders of magnitude lower than observed. Similar or greater discrepancies are found for the other experimental conditions.

[54] A full sensitivity analysis is performed on the model in order to determine the nature of the feedback expected by perturbations to model parameters, e.g., reaction rate constants, with respect to Cl_2 formation. A full Jacobian matrix indicating change in concentrations versus change in parameters was generated for the system. Table 8 shows the response normalized as a percent change of the Cl_2 concentration due to a +1% change in the corresponding parameter. The results of the sensitivity analysis provide much insight while bearing in mind the caveat that any *local derivative*, such as those calculated here, should not be wildly extrapolated.

[55] The Cl_2 formation due to known chemical reactions is explained by an alternative yet similar mechanism to the one proposed earlier by *Oum et al.* [1998],



with the critical steps given by the formation of the $\text{HOCl}^{\bullet-}$ radical ion that may then react with the abundant chloride ions to form the chlorine radical ion, $\text{Cl}_2^{\bullet-}$. In the model, the pH of the droplets begins at equilibrium with the chamber CO_2 at a value near 6.6 and quickly reaches a value of 8.7, tapering to 9 by the end of the experimental simulation, due to the formation of hydroxide ions in (35). Consequently,

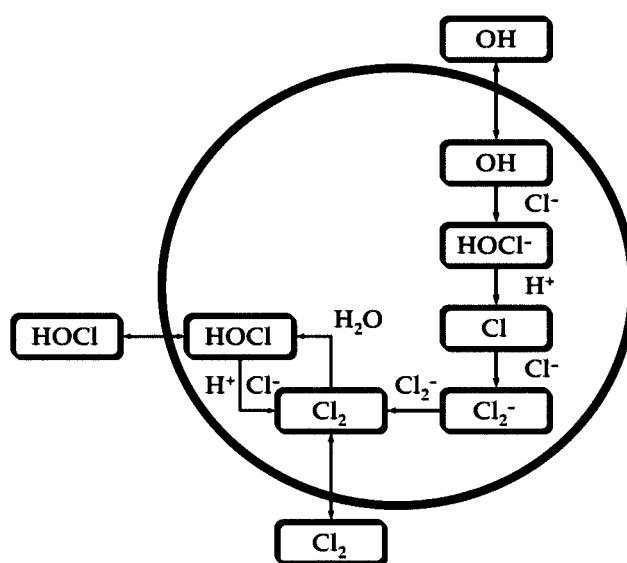


Figure 1. Graphical representation of known aqueous-phase chemical reactions potentially responsible for observed Cl₂.

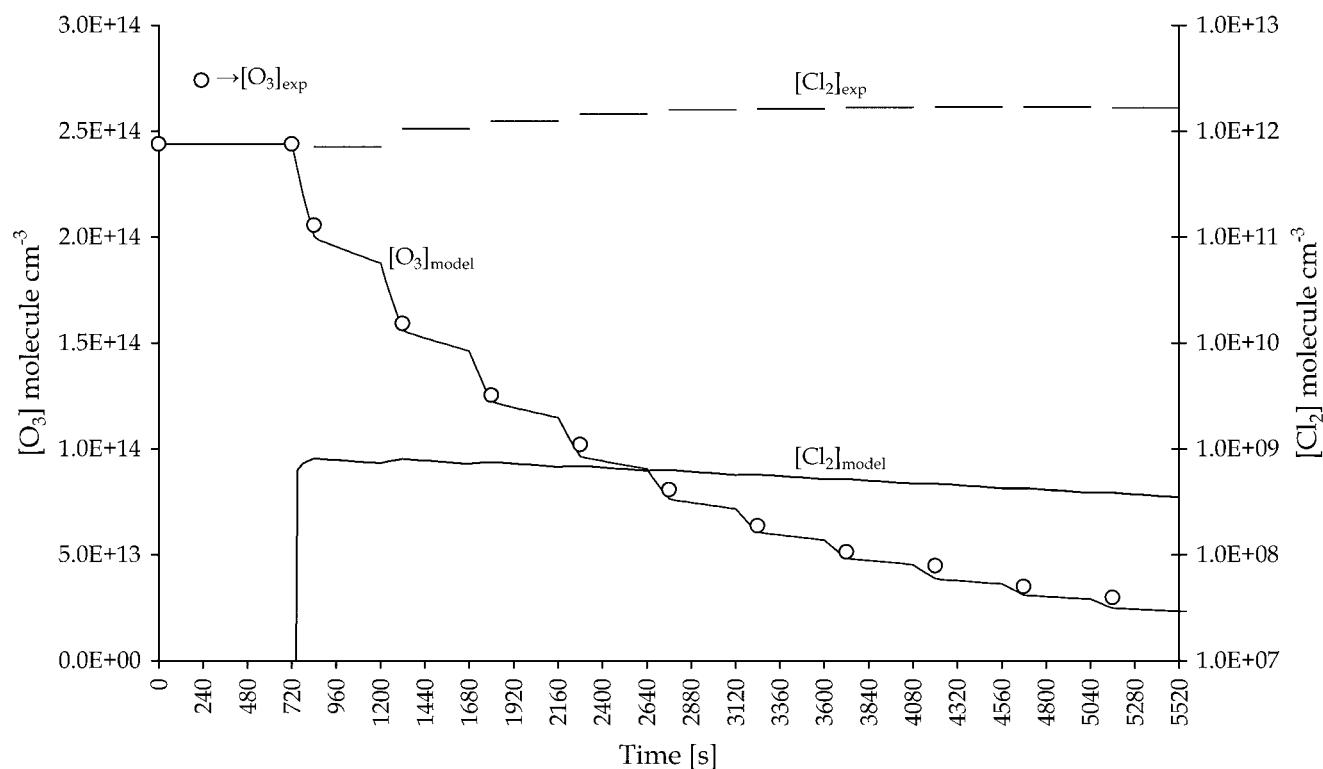


Figure 2. Results of aerosol chamber modeling (middle O₃ experiment) using all known aqueous-phase and gas-phase chemical reactions (base case scenario). Cl₂ concentrations are shown on a logarithmic scale to better illustrate the discrepancy between observations and predictions using known chemistry.

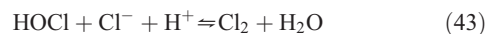
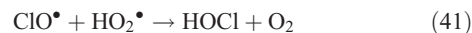
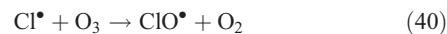
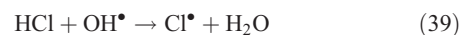
the reaction of the HOCl^{•-} intermediate with H⁺ is progressively slower compared to the reaction with Cl⁻ and throughout the simulation.

[56] The decay in the Cl₂ concentration is attributable to the combined effect of reaction with OH[•] in the gas phase, photolysis, loss to the walls of the experimental chamber, and dilution during API-MS sampling. The most important, albeit of lesser overall impact, loss routes in the aqueous phase are reaction with OH⁻ and HO₂⁻. HO_x radicals are removed from the system forming O₂ and H₂O mainly by gas-phase reactions resulting in a net removal of radicals or potential radical-forming species: O₃ + OH[•], OH[•] + HO₂[•], H₂O₂ + OH[•] and O₃ + HO₂[•].

[57] The confidence level for the rate constant values of the most significant reactions leading to chlorine formation is high, particularly for the gas-phase reactions. Some mass transfer parameters are not as well known but their uncertainty still cannot justify the larger than three order-of-magnitude discrepancy between the model results and experimental observations. For instance, variation of the OH[•] and HO₂[•] accommodation coefficients within their range of uncertainty, 0.01–1.0 [Hanson *et al.*, 1992; Takami *et al.*, 1998], results in little appreciable Cl₂ increase compared to the chamber measurements. Furthermore, any realistic error regarding the aerosol population, initial concentrations or model formulation cannot result in a 1000-fold increase in chlorine concentrations.

[58] Is there an unknown chemical reaction or physical process that could account for the Cl₂ formation? In view that the droplet pH is a paramount factor in the production

of chlorine, a series of simulations are executed in order to evaluate the effect of acidity. A series of model simulations specifying a constant droplet pH starting with five and lowering to one are performed. Figure 3 shows that only at a constant droplet pH of 1.5 do model results match experimental observations. Sensitivity studies (Table 8) and analysis of this simulation show that the chlorine formation process that prevailed in the base case simulation is overwhelmed by a chlorine formation route initiated by hydrochloric acid displacement,



followed by a chain of reactions in the gas phase forming hypochlorous acid. HOCl is subsequently taken up by the droplets and forms Cl₂ via the reverse hydrolysis reaction.

[59] There is no apparent reason why the droplets would become so acidic, and a constant pH is even less sensible. Nonetheless, two possibilities are examined. First, a series of simulations explore the effect of lowering the initial pH

Table 8. Local Sensitivity of Final Cl₂ Concentrations to Model Parameters

Parameter	Simulation		
	Base case	pH = 1.5	Interface
<i>LWC</i>	0.99%	0.99%	−0.11%
<i>Surface area</i>			0.54%
<i>R (mass transfer)</i>	−0.60%	−0.70%	0.03%
<i>R (aq. diffusion)</i>		−0.01%	0.32%
[NaCl] ₀	−1.74%	−0.45%	1.24%
[O ₃] ₀	0.03%	0.21%	0.19%
[CO ₂] ₀	0.22%		0.08%
<i>γ_s</i>			0.54%
<i>RH</i>	0.17%	0.28%	0.13%
<i>Dilution factor</i>	9.19%	5.94%	5.53%
<i>k₂</i>	O ₃ + hν → O(³ P) + O ₂	0.20%	0.35%
<i>k₅</i>	Cl ₂ + hν → Cl• + Cl•	−0.04%	−0.03%
<i>k₉</i>	O(¹ D) + O ₂ → O(³ P) + O ₂	−0.05%	−0.04%
<i>k₁₂</i>	O(¹ D) + H ₂ O → OH• + OH•	0.17%	0.12%
<i>k₁₃</i>	O(¹ D) + N ₂ → O(³ P) + N ₂	−0.12%	−0.09%
<i>k₁₈</i>	O ₃ + OH• → HO ₂ • + O ₂	−0.23%	−0.26%
<i>k₂₁</i>	OH• + HO ₂ • → O ₂ + H ₂ O	*	−0.05%
<i>k₂₃</i>	O ₃ + HO ₂ • → OH• + O ₂ + O ₂	*	−0.02%
<i>k₃₀</i>	OH• + Cl ₂ → Cl• + HOCl	−0.22%	−0.15%
<i>k₃₁</i>	OH• + ClO• → HCl + O ₂		
<i>k₃₂</i>	OH• + ClO• → HO ₂ • + Cl•	−0.06%	
<i>k₃₄</i>	OH• + HCl → Cl• + H ₂ O	−0.07%	−0.03%
<i>k₃₅</i>	OH• + HOCl → ClO• + H ₂ O	0.24%	
<i>k₃₈</i>	HO ₂ • + ClO• → HOCl + O ₂	−0.07%	
<i>k₆₀</i>	Cl ₂ + Wall → Loss	0.12%	0.02%
<i>k₁₀₀₉</i>	OH• + HO ₂ • → HO ₂ • + OH•	−0.33%	−0.15%
<i>k₁₀₃₆</i>	O ₃ + HO ₂ • → HO ₂ • + O ₂	*	*
<i>k₁₁₀₈</i>	OH• + Cl• → HOCl•	0.04%	0.03%
<i>k₁₁₀₉</i>	HOCl• → OH• + Cl•	−0.03%	−0.02%
<i>k₁₁₁₄</i>	HOCl• + H• → HOClH•		0.02%
<i>k₁₁₁₆</i>	Cl ₂ • + Cl ₂ • → Cl ₂ + 2 Cl•	0.25%	0.03%
<i>k₁₁₁₈</i>	HOCl• + Cl• → HOCl• + Cl•	0.03%	*
<i>k₁₁₁₉</i>	Cl ₂ • + HO ₂ • → 2 Cl• + H• + O ₂		−0.06%
<i>k₁₁₂₀</i>	Cl ₂ • + O ₂ • → 2 Cl• + O ₂	*	*
<i>k₁₁₃₃</i>	HOCl + Cl• + H• → Cl ₂ + H ₂ O		0.35%
<i>k₁₁₃₄</i>	Cl ₂ + OH• → HOCl + Cl•	−0.02%	−0.15%
<i>k₁₁₃₇</i>	Cl ₂ + HO ₂ • → 2 Cl• + H• + O ₂	*	−0.10%*
<i>k₁₁₃₈</i>	HOCl + HO ₂ • → Cl• + O ₂ + H ₂ O	0.02%	0.01%
<i>k₁₂₁₃</i>	CO ₃ • + Cl ₂ • → CO ₃ • + 2 Cl•	−0.50%	
<i>k₁₂₁₅</i>	Cl ₂ + CO ₃ • + H ₂ O → HOCl + Cl• + HCO ₃ •		−0.12%
<i>α_{OH}</i>		0.33%	0.38%
<i>α_{HO₂}</i>			−0.03%
<i>H_{O₃}</i>			0.02%
<i>H_{OH}</i>		0.03%	*
<i>H_{HO₂}</i>		−0.07%	*
<i>H_{H₂O₂}</i>	*	*	−0.08%*
<i>H_{HCl}</i>		−0.23%	
<i>H_{HOCl}</i>	*	0.36%	0.01%
<i>H_{Cl₂}</i>	*		−0.43%
<i>H_{CO₂}</i>	0.22%		0.08%
<i>K_{H₂O}</i>	*		−0.15%
<i>K_{H₂O₂}</i>	*		−0.08%*
<i>K_{HO₂}</i>		*	
<i>K_{HCl}</i>		−0.23%	
<i>K_{CO₂}</i>	0.22%		0.08%
<i>K_{HCO₃•}</i>	−0.24%		0.07%

Sensitivity shown when the absolute value of the normalized Jacobian terms, $J_{i,j}$, are greater than 0.01%, i.e., $|(\partial C_i / \partial P_j) \times (\partial P_j / C_{i, \text{test}})| > 0.01\%$. For the above table, C_i is the final $[\text{Cl}_2]$, $C_{i, \text{test}}$ is the final $[\text{Cl}_2]$ under the normal simulation, and P_j represents the parameters tested with $\partial P_j = 0.01$ Pj. Sensitivities of Cl₂ concentrations to the above parameters are invariably higher during the earlier portions of the experimental simulations. Asterisks (*) denote moderately appreciable Cl₂ Sensitivity, normalized $|J_{[\text{Cl}_2], j}| > 0.01\%$, found during earlier periods of the simulation. Early moderate sensitivity is also found in general for reactions gas-phase reactions 20, 22, 24, and 25. Maximum values shown in italics for sensitivity to three H₂O₂/HO₂-related parameters in interfacial mechanism simulation. An increase in the dilution factor equals a reduction in overall dilution.

of the droplets beyond expected values. This would account for the incidence of an unknown phenomenon that could potentially bring the droplets to a lower initial pH, such as sequestering of sodium cations during particle generation or some unexplained generation of hydrochloric acid. Figure 4

shows that in the scenario that matches the initial rate of formation, the droplets have lost 25% of their sodium cations. This is tantamount to having an initial hydrochloric acid concentration of 1.08 M and pH of −0.4, quickly equilibrated during the preliminary 12-min dark period to a

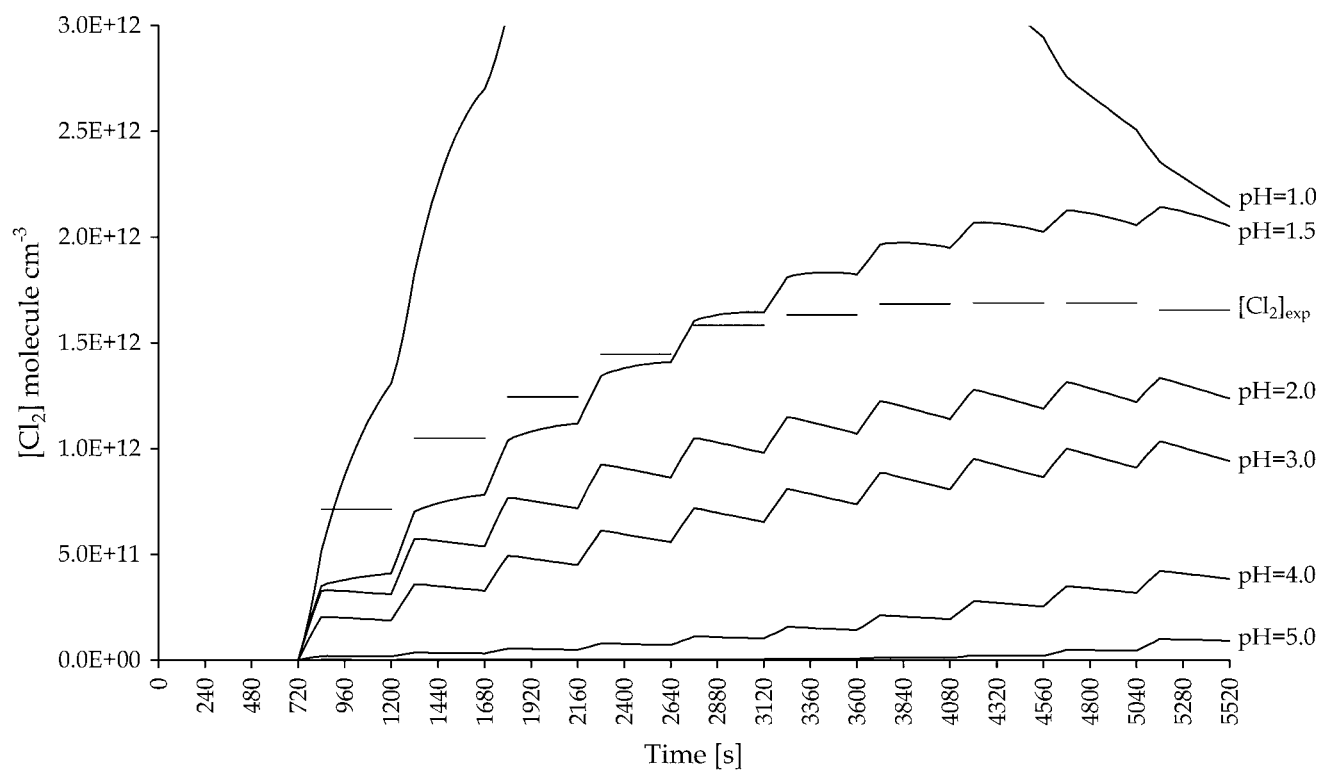


Figure 3. Results of aerosol chamber modeling (middle O_3 experiment) at fixed pH values. The ozone time series are all similar to the ozone time series shown on Figure 2.

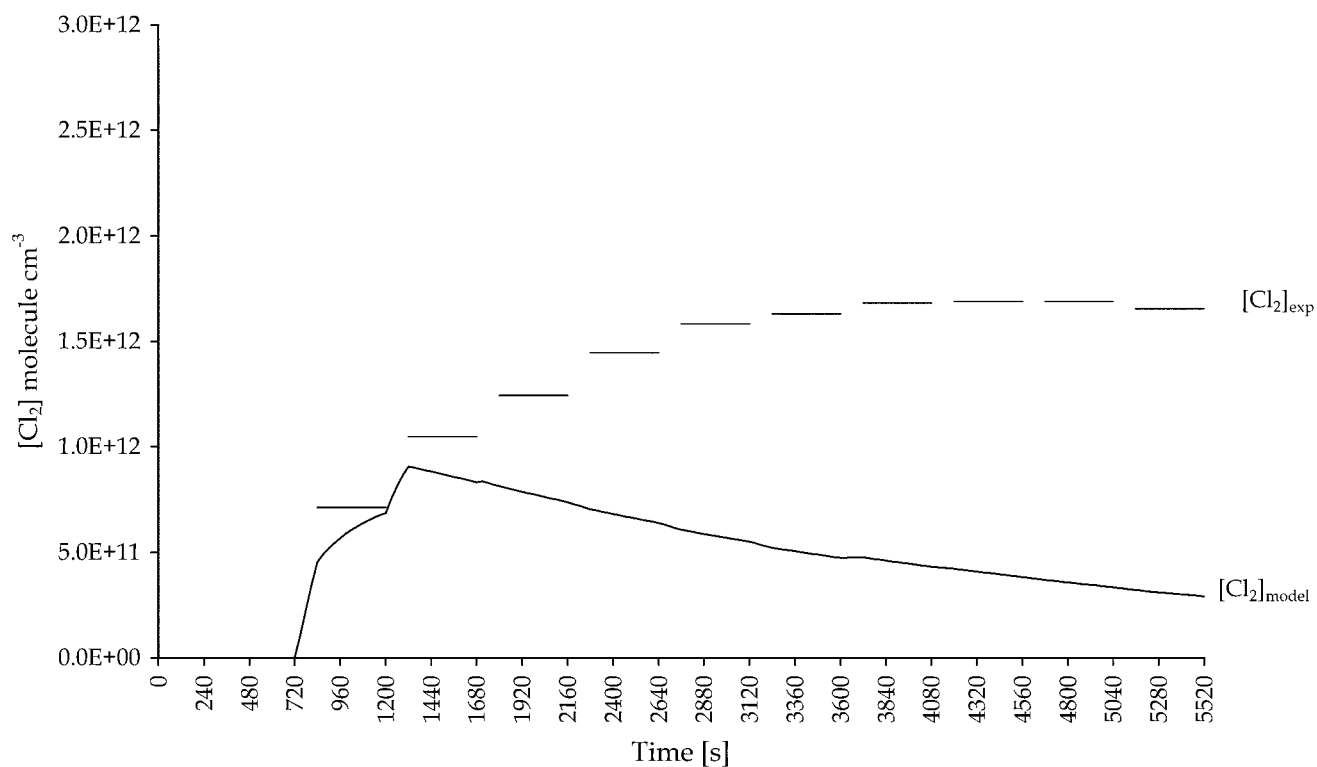


Figure 4. Results of aerosol chamber modeling where 25% initial NaCl substituted by HCl . The ozone time series is similar to the ozone time series shown on Figure 2.

pH of 0.88. Not only is this case highly improbable, it also fails to reproduce the data adequately.

[60] An additional modeling run is performed adding the dissociation of HOCl^{•-}, i.e.,



to provide a channel for acidification of the particles. Jayson *et al.* [1973] propose such a reaction but the model demonstrates that a considerable rate of Cl₂ formation is obtained only if HOCl^{•-} is found an infinitely strong acid, i.e., $\text{p}K_a \ll 0$. This is highly improbable given that the $\text{p}K_a$ of HOClH[•] is ~ 2.5 ; the $\text{p}K_a$ of HOCl^{•-} should be indicative of the second deprotonation, thus higher than the $\text{p}K_a$ of HOClH[•] and accordingly much higher than zero.

[61] Finally, consideration of chamber wall reactions as the source of the molecular chlorine is safely eliminated. Oum *et al.* [1998] detail the precautionary measures taken in order to minimize any effect of the chamber walls during the experiments and exclusion of the walls as a potential source. Furthermore, the chamber walls are found to be considerable sinks of Cl₂ and HOCl. These effects are modeled as irreversible losses reflecting the current knowledge of the system.

3.2. Interfacial Chemical Mechanisms

[62] Clearly, the observed Cl₂ formation cannot be reconciled by established chemical mechanisms—the disagreement between the experimental and model results is entirely too large. However, as expressed earlier, the gas–liquid interface does provide a sound platform for hitherto investigated chemical reactions to occur between a reactive gas-phase species and available surface chloride.

[63] Two observations are presented before proceeding. First, the values of the gas-phase concentrations of O₃, HO₂[•] and OH[•] are not affected significantly by reactions in the aqueous phase or with liberated chlorine species throughout the span of the experimental simulations—an assumption found correct by a posteriori analyses. This assumption is not satisfied for H₂O₂ in any model simulation including surface chemistry mechanisms as described below.

[64] Second, the determination of the surface area and volume of the aerosol population is subject to significant error. In all three of the experiments presented in this paper, careful attempts are made to introduce equivalent aerosol populations using the same laboratory procedures. The aerosol population is counted by a condensation particle counter and measured by a differential mobility analyzer (DMA) twice before and twice after the experiment. In order to minimize error in volume and surface area calculations, each output from the DMA (cutoff radius $\sim 0.5 \mu\text{m}$) is fit to a lognormal distribution, integrated over all droplet radii and an average over the four measurements is obtained due to the complexities associated with particle coagulation, loss during API-MS sampling and condensation to chamber walls. There is a +60% difference in aerosol volume and +85% difference in surface area when comparing the low ozone and high ozone experiments. Even though the variations appear genuine, the extent to which these differences could be attributable to errors in aerosol measurement or to the lognormal assumption is unclear—at best, the four-sample averaged lognormal fit can only estimate and not

affirm the dimensions of the particles in the region $R_p \geq 0.5 \mu\text{m}$, and it does not consider particle volume adhered to the chamber walls. These possible “errors” cannot justify the thousand fold discrepancy in Cl₂ concentrations in base case model simulations, but should be considered when evaluating the suitability of a potential surface mechanism. In all simulations, the values for surface area and volume, as well as the radius used in the mass transfer and aqueous-phase diffusion equations, are determined using the four sample averaged log normal fit parameters.

[65] The initial rate of Cl₂ formation, calculated from the concentration after 2 min of photolysis, if resultant from a surface reaction is expected adhere to an expression given by:

$$R = \gamma_s \frac{\bar{c}}{4} A[G]. \quad (45)$$

where [G] is the average concentration of the reacting species and all other variables are as defined in section 2.7. For all experiments, estimates of the initial Cl₂ production rate fall in the narrow range of $4.2\text{--}5.9 \times 10^9 \text{ molecule cm}^{-3} \text{ s}^{-1}$ correlating better with the prediction of a narrow range of initial (averaged over the first 2 min) OH[•] concentrations, $1.1\text{--}1.3 \times 10^{10} \text{ molecule cm}^{-3}$, than with the prediction of much wider ranges of initial [HO₂[•]], $2.3\text{--}8.1 \times 10^{10} \text{ molecule cm}^{-3}$, and initial [O₃], $0.6\text{--}3.0 \times 10^{14} \text{ molecule cm}^{-3}$. The middle and high initial O₃ experiments showed remarkably similar Cl₂ formation rates (within 3%) even though aerosol measurements yield 40% more surface area in the latter experiment. Considering the product of gas concentration and surface area, the results are still more consistent with OH[•] as the reactive species both considering and neglecting the potential errors of aerosol measurement.

[66] The overall surface mechanism proposed is equation (30). The overall enthalpy change of the reaction is calculated at $-230.6 \text{ kJ mol}^{-1}$ [Lide, 1998]; the proposed surface mechanism is exothermic and thermodynamically favorable. The formation of a chloride–hydroxyl complex is assumed to be the rate determining step of the overall reaction (30); the possible fates of the OH●●●Cl^{•-}_{surface} species are discussed subsequently. The overall rate of Cl₂ formation is calculated according to equations (28) and (29) from section 2.7. An optimal value of 2 for γ_s' is obtained by fitting Cl₂ predictions to the results of the middle ozone experiments.

[67] Figures 5, 6, and 7 illustrate that for all experiments, including the higher and lower initial ozone concentration experiments, the correspondence of the model results to the experimental data (given the level of uncertainties inherent to the model and in the experiments) is quite good; the modeled Cl₂ concentration profiles are essentially correct with maximum deviations bounded by 2/3 and 1 2/3 multiples of the experimental data. Regardless of the details of the elementary reactions, the model results support the theory of a surface-enhanced Cl₂ generation mechanism.

[68] The present authors in joint investigation with theoretical and experimental chemical researchers [Knipping *et al.*, 2000] presented as a primary hypothesis that the product of the initial reaction between the approaching gas-phase hydroxyl radical and the exposed surface chloride ions is a relatively stable surface complex, OH●●●Cl^{•-}_{surface} (equation (27)). No experiments have proven the existence of OH●●●Cl^{•-}_{surface}; this postulate is offered above all to inspire

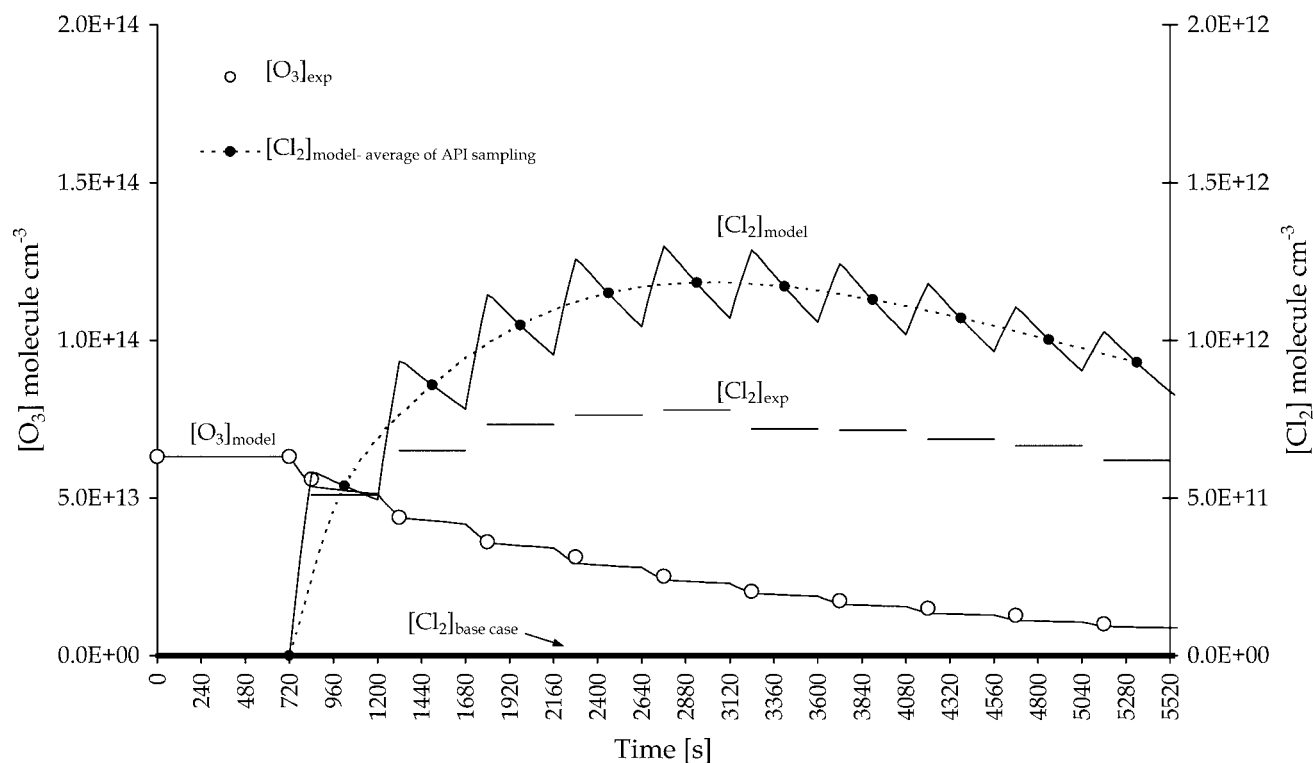


Figure 5. Experimental and model Cl_2 time series using proposed interfacial mechanism: low ozone experiment.

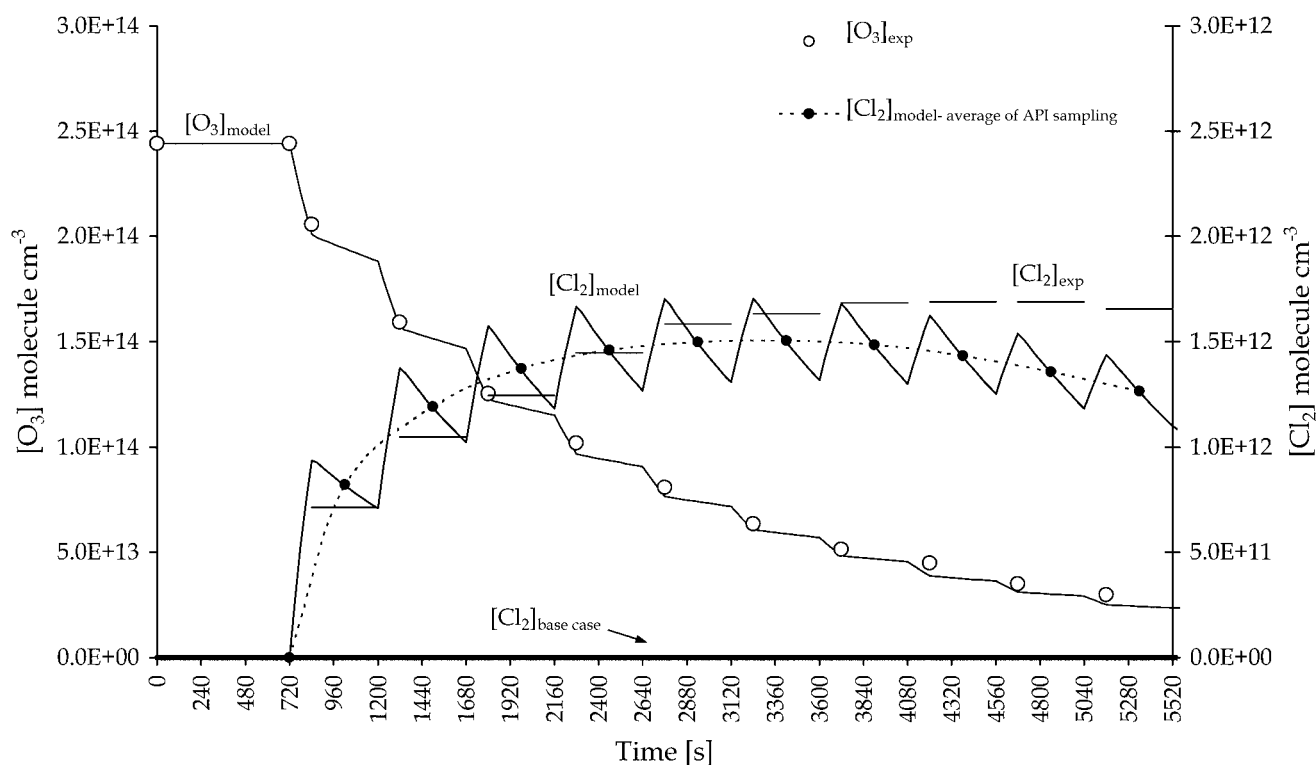


Figure 6. Experimental and model Cl_2 time series using proposed interfacial mechanism: middle ozone experiment.

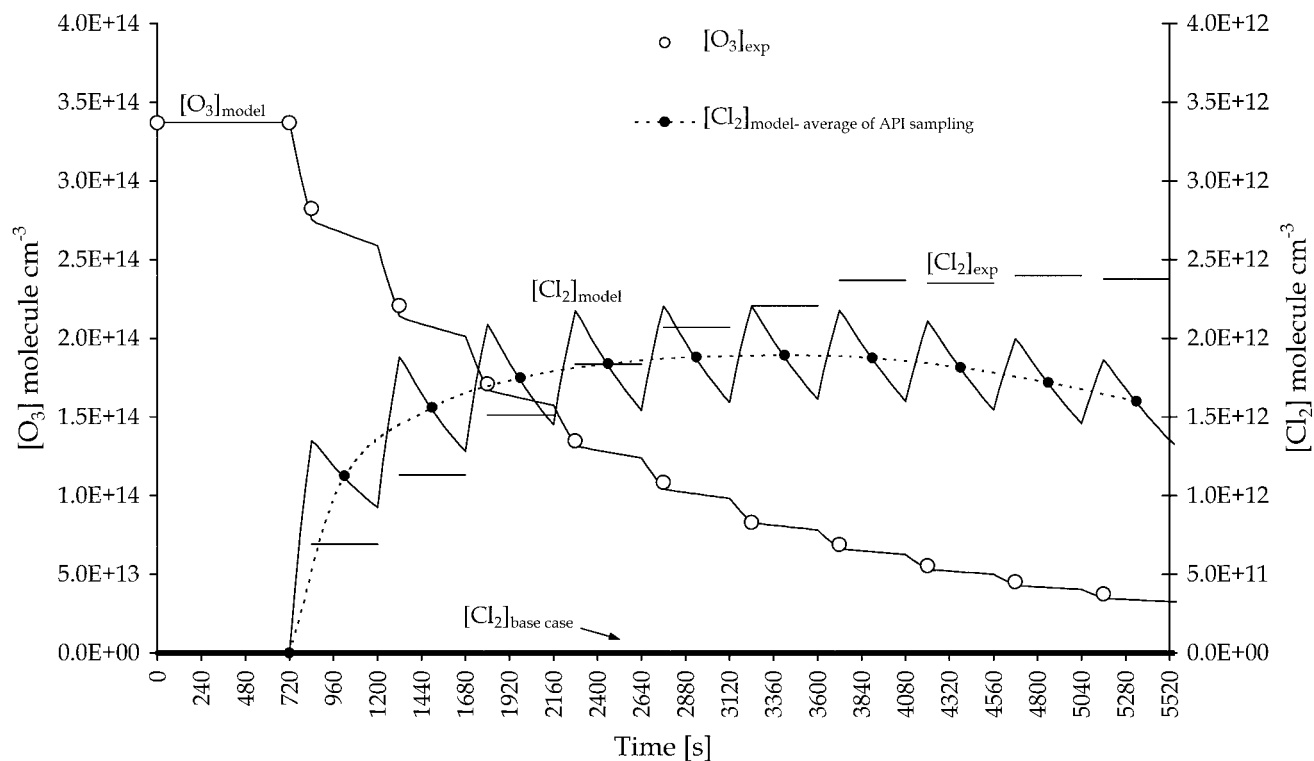
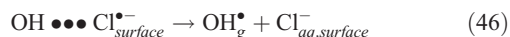


Figure 7. Experimental and model Cl₂ time series using proposed interfacial mechanism: high ozone experiment.

the search for evidence/counterevidence and arouse insightful feedback. Still, the adequacy of the overall rate expression is not dependent on the existence of the chloride–hydroxyl complex even though the basis is motivated by the formation of such a species.

[69] Two fates of the proposed surface chloride–hydroxyl radical ion have been considered. In both cases the bound complex, $\text{OH}\bullet\bullet\bullet\text{Cl}_{\text{surface}}^-$, is assumed to remain in a 1 nm shell ($\sim 6 \times$ the ionic radius of Cl^- , 1.81 Å) on the exterior of the droplet, as suggested by density profiles of sodium, chloride and oxygen atoms in water slabs [Jungwirth and Tobias, 2000]. The species should be only somewhat stable and therefore undergo decomposition back to reactants



[70] The first prospective fate of the bound complex, illustrated in Figure 8, is the self-reaction



Due to similarity with the $\text{Cl}_2^{\bullet-}$ radical ion, the rate constant of reaction 47 can be estimated from $\text{Cl}_2^{\bullet-}$ recombination at $1.8 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$ [Jacobi et al., 1999]. Alternatively, $\text{OH}\bullet\bullet\bullet\text{Cl}_{\text{surface}}^-$ may react directly with surface chloride ions to form $\text{Cl}_2^{\bullet-}$ and OH^- . The rate constant for this reaction is matched to the rate constant for the reaction of $\text{HOCl}^{\bullet-}$ and Cl^- , i.e., $1 \times 10^4 \text{ M}^{-1} \text{ s}^{-1}$ [Grigor'ev et al., 1987]. The $\text{Cl}_2^{\bullet-}$ radical anion is assumed to stabilize at the interface where it may build-up in concentration and self-react to form Cl_2 and two chloride

ions, or react with another $\text{OH}\bullet\bullet\bullet\text{Cl}_{\text{surface}}^-$ to form Cl_2 and separate chloride and hydroxide ions; preliminary molecular dynamic simulations indicate that $\text{Cl}_2^{\bullet-}$, once formed, could indeed remain anchored at the surface of concentrated aqueous particles (P. Jungwirth, personal communication). In both cases, the rate of $\text{OH}\bullet\bullet\bullet\text{Cl}_{\text{surface}}^-$ decomposition must be regulated to $\sim 10^4 \text{ s}^{-1}$ in order to guarantee formation of $\text{OH}\bullet\bullet\bullet\text{Cl}_{\text{surface}}^-$ as the rate-limiting step in Cl_2 generation.

[71] An unbiased assessment of the overall mechanism of chlorine production (30) leads to the proposal of a debatably more straightforward product during the reaction of $\text{OH}^\bullet$ and Cl^- , the $\text{HOCl}^{\bullet-}$ species originally reported by Jayson et al. [1973]. The differences between the surface complex $\text{OH}\bullet\bullet\bullet\text{Cl}_{\text{surface}}^-$ and the bulk species $\text{HOCl}^{\bullet-}$ may or may not be subtle and further research is warranted. Jayson et al. [1973] calculated an absorption spectrum for a species in solution they denoted as $\text{HOCl}^{\bullet-}$. Aqueous-phase radical chemistry mechanisms based on this species are still undergoing much investigation [McElroy, 1990; Buxton et al., 1998; Alegre et al., 2000]. However, in studies of the transition state of the reaction



Davis et al. [1994] concluded that an OHCl^- ion formed in the gas phase by electron impact of a mixture of HCl and N_2O is best represented as $\text{OH}\bullet\bullet\bullet\text{Cl}^-$ and calculated a photodetachment spectrum for the species. Electron spin resonance (ESR) studies of Cl^- in frozen aqueous solution and molecular orbital theory calculations by Sevilla et al. [1997] support the existence of a $\text{Cl}\cdots\text{H}_2\text{O}$ adduct in solution

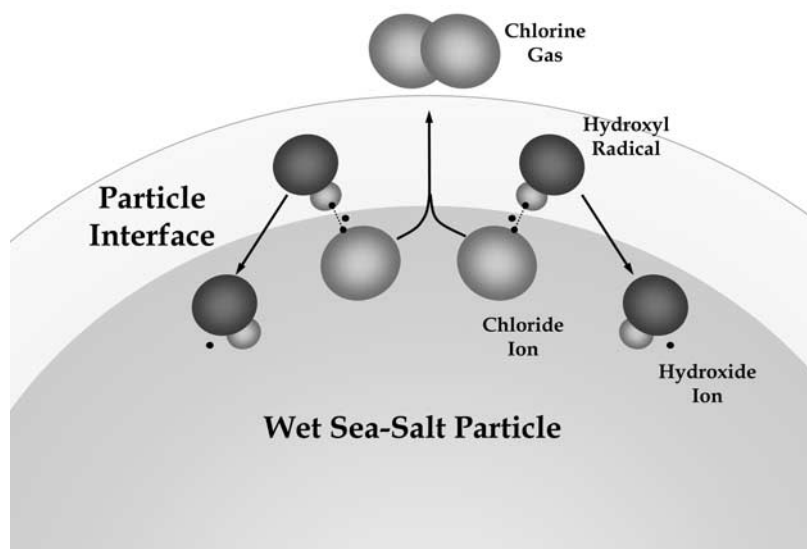


Figure 8. Potential mechanism for Cl_2 production.

rather than $\text{Cl}-\text{OH}^-$. These studies showed ion–dipole interactions between the hydrogen of $\text{OH}^\bullet$ and chloride ion could generate an $\text{OH}\cdots\text{Cl}^-$ complex, but it also determined that the bond strength of a $\text{Cl}^-\cdots\text{HO}(\text{H}_2\text{O})_n$ structure decreased as the number of hydrating waters, n , increased. The evidence suggests that the presence of an $\text{OH}\cdots\text{Cl}^-$ complex is likely only at gas–liquid interfaces, while the structure of the bulk $\text{HOCl}^\bullet$ species could be better represented as $\text{Cl}(\text{H}_2\text{O})-\text{OH}^\bullet$. At any rate, model simulations performed where $\text{HOCl}^\bullet$ is produced at the interface, followed by diffusion and reaction in the bulk of the droplet, predict low Cl_2 levels similar to the results of the base case simulation.

[72] Other fates of the $\text{OH}\cdots\text{Cl}_{\text{surface}}^\bullet$ complex or surface anchored $\text{Cl}_2^\bullet$ have been explored without success (e.g., if Cl_2 is generated by electron photodetachment of $\text{Cl}_2^\bullet$ within the droplet, using $\sigma = 200 \times 10^{-20} \text{ cm}^2$ at 254 nm [Adams *et al.*, 1995], significant induction times and incorrect peak $[\text{Cl}_2]$ values are predicted). The viability of the principal mechanisms presented here underscores the importance of further investigating concentrated salt aerosol interfaces.

3.3. Sensitivity Studies of the Overall Interfacial Mechanisms

[73] The sensitivities of Cl_2 formation with respect to model parameters for simulations incorporating the overall interfacial mechanism (30) and kinetic rate expression (28) are presented as normalized Jacobian terms shown on Table 8. There is no evidence that the observed Cl_2 formation in these model simulations can be drastically (on the scale of orders of magnitude) altered by small perturbations of the listed parameters, but the sensitivity results, together with a thorough inspection of the simulation output, do indicate that there are several parameters that could significantly affect the intensity and temporal behavior of the modeled Cl_2 production.

[74] The principal source of Cl_2 in these simulations is the reaction at the interface. This interfacial mechanism

produces hydroxide ions in solution thus increasing pH. After the first 2 min of photolysis, the model predicts that particle pH levels rise to ~ 10.5 and gradually reach a value near 10.8 at the end of the simulation. Any contributions from the reactions of HOCl with Cl^- , in the presence and absence of H^+ , as well as the recombination of $\text{Cl}_2^\bullet$ in bulk solution are negligible. Due to the increased alkalinity of the droplets, Cl_2 decomposition reactions with OH^- , HO_2^- and CO_3^{2-} become major loss routes for Cl_2 . The rate of chlorine loss due to these aqueous-phase reactions is greater than the effect of dilution and wall loss, but the importance of the latter pair is heightened by their ability remove available chlorine from the modeled system. The sensitivity results show that as total particle surface area is increased, the final levels of Cl_2 increase due to the enhancement of the surface reaction. However, as the liquid water content of the aerosol increases, final Cl_2 levels decrease due to the efficient loss processes for dissolved chlorine in alkaline aqueous solution. These two results alone indicate that optimal Cl_2 is likely when the surface to volume ratio is maximized, i.e., smaller particles. (As expected, in simulations where the pH of the particles remains low and the interfacial mechanism is excluded, an increase of aerosol volume increases the final $[\text{Cl}_2]$ in an essentially linear fashion.)

[75] The sensitivity of Cl_2 formation to particle size, independent of overall volume and surface area, is intriguing. The chlorine production is insensitive to the radius of the particles used in determining the mass transfer terms, k_{mt} , but is moderately sensitive to the radius of the particles used in determining the spatially averaged aqueous-phase reactions rates, i.e., the radius used in determining the extent of the aqueous-phase diffusion limitation. In the current model, these two radii are equivalent, but the sensitivities are determined separately in order to distinguish between the two limitations to mass transfer. Whereas the sensitivity to the liquid water content and surface area indicates that smaller particles, i.e., large surface to volume ratios, are most favorable to chlorine production, the sensitivity to the

mass transfer and aqueous-phase diffusion limitations indicate that larger particles are better Cl_2 generators.

[76] These results suggest the existence in the chamber, and possibly the atmosphere, of an ideal particle size for Cl_2 activation; however, the value of this optimal size is not explored. (The equations interrelating relative humidity to particle size and salt concentration [Pruppacher and Klett, 1997] are complex. A study to determine how the singular effects contributing and limiting Cl_2 production change with increasing particle size would yield an optimization problem with many interdependent parameters.) The sensitivity studies presented here are performed with all variables independent of each other. For example, both the droplet radius (and thus surface area and volume) and the initial sodium chloride concentration are expected to depend on relative humidity, yet all these parameters were explored individually. The advantage of this uncoupled analysis is manifest in the sensitivity (recalling that the sensitivity is a local derivative calculated at normal unperturbed conditions) to initial $[\text{NaCl}]$: in the base case, an increase in initial concentration causes a decrease in final $[\text{Cl}_2]$ due to the increase in ionic strength and the consequent kinetic salt effect enhancement of k_{1213} (which is much more pronounced than the enhancement on k_{1116}); while in the constant pH simulation, an increase in ionic strength causes a decrease in final $[\text{Cl}_2]$ due to a kinetic salt effect interference on k_{1133} ; in both simulations, variations in the ionic strength also cause modifications in the calculation of effective Henry's law constants and acid/base equilibrium constants that ultimately result in an impediment to Cl_2 formation. On the other hand, in the interfacial mechanism, higher initial concentrations of sodium chloride are very effective in generating an increase in Cl_2 due to the amplification of the surface reaction.

[77] Indeed, the interfacial sensitivity studies show that higher salt concentrations are more effective in increasing chlorine production than are higher concentrations of O_3 or CO_2 . Nonetheless, the effect of CO_2 is rather considerable in the system. While the interfacial mechanism does not require a contribution of H^+ for Cl_2 production, the chlorine decomposition rate, as discussed earlier, is highly dependent on the droplet pH. The presence of carbon dioxide provides sufficient buffering and keeps modeled pH values in the range of 10.5–10.8, whereas if no CO_2 were present the droplets would reach modeled pH values of ~ 13 and Cl_2 production would be greatly restricted. (Although the presence of CO_3^{2-} does add another aqueous-phase loss for Cl_2 , lowering the effectiveness of the reaction with hydroxide ion is much more effective for maintaining higher Cl_2 values.) CO_2 is present in the chamber at concentrations near 17 ppm. Model evaluations determine that Cl_2 production due to the proposed mechanism is insensitive to the exact value of $[\text{CO}_2]$ when initial concentrations are above 10 ppm. However, a significant drop in peak $[\text{Cl}_2]$ is revealed for lesser initial carbon dioxide concentrations, and a sharp decline is evident when $[\text{CO}_2]_0$ decreases below ~ 1 ppm. The role of CO_2 with respect to chlorine generation has been indicated recently in studies of the thermal decomposition of ozone in sodium chloride solutions [Levanov et al., 2000].

[78] In the kinetic expression of the overall interfacial mechanism, the magnitude of the free variable, γ_s' , is

adjusted in order to obtain an optimal match to the production rate and peak values of chlorine in the laboratory chamber. Model evaluations yield an optimum value $\gamma_s' = 2$. Although it is not immediately intuitive that γ_s' should possess a value greater than unity, there are arguments in support of higher values. (However, the condition $\gamma_s \leq 1$ must be constantly upheld.) Practically converged quantum chemical calculations suggest that surface chloride ions in effect attract approaching hydroxyl radicals [Knipping et al., 2000], thus enhancing the rate of scavenging. Such an enhancement due to attractive forces is analogous to the increase of reaction rates beyond the predicted aqueous-phase diffusion-controlled limit exhibited by proton transfer reactions [Benson, 1960]. Furthermore, there are indications that $\text{OH}^\bullet$ radicals could “roll” on aqueous surfaces. As a result, γ_s' may have a functional form different than a simple constant value and an alternative model for the overall interfacial mechanism rate could be better suited. The sensitivity of Cl_2 levels to γ_s' is high and further studies should be performed in order to better determine the character of this variable.

[79] Figure 9 shows time series depicting the rate of Cl_2 production and loss corresponding to the major chemical and physical processes. Although the rate of reaction of Cl_2 with $\text{OH}^\bullet$ in the gas phase is considerable, the formation of chlorine via the proposed interfacial mechanism is an order of magnitude higher resulting in an overall net production effect for $\text{OH}^\bullet$. The effects of experimental processes such as the small amount of Cl_2 photolysis, irreversible losses to the walls and dilution during API-MS sampling, are quite significant. (The uncertainty of the dilution factor is much less than one percent thus lowering the impact of its high sensitivity.) Nonetheless, the rate of Cl_2 production is predominated by production at the interface and losses in the aqueous phase. The uncertainties associated with the rate constants for the three main alkaline Cl_2 decomposition processes—reaction with OH^- , HO_2^- and CO_3^{2-} —are rather high, amplifying their sensitivity and hence their importance in developing mathematical chemistry models capable of accurately reproducing Cl_2 concentrations for variable conditions. Due to the importance and uncertainty of these alkaline chlorine decomposition reactions, a brief review of their study and the basis for selecting rate constants is provided in the following sections. In addition, Table 9 illustrates the sensitivity of peak and final model Cl_2 concentration values, expressed as the average concentration measured during the corresponding API dark sampling period, to order-of-magnitude changes in the selected rate constants for these alkaline Cl_2 decomposition processes within the limits reported in the scientific literature.

[80] The uncertainties regarding the geometric parameters (e.g., surface area and liquid water content) of the experimental aerosol populations were discussed in the previous section. Together with the results of these sensitivity analyses, especially given the uncertainties with respect to alkaline Cl_2 decomposition processes, the discrepancies between the lower and higher ozone model results and the experimental data can be explained by variations in model input and/or parameters. These sensitivity studies conclude a methodological analysis of the chemical system. The complete analysis strongly corroborates the suggestion of

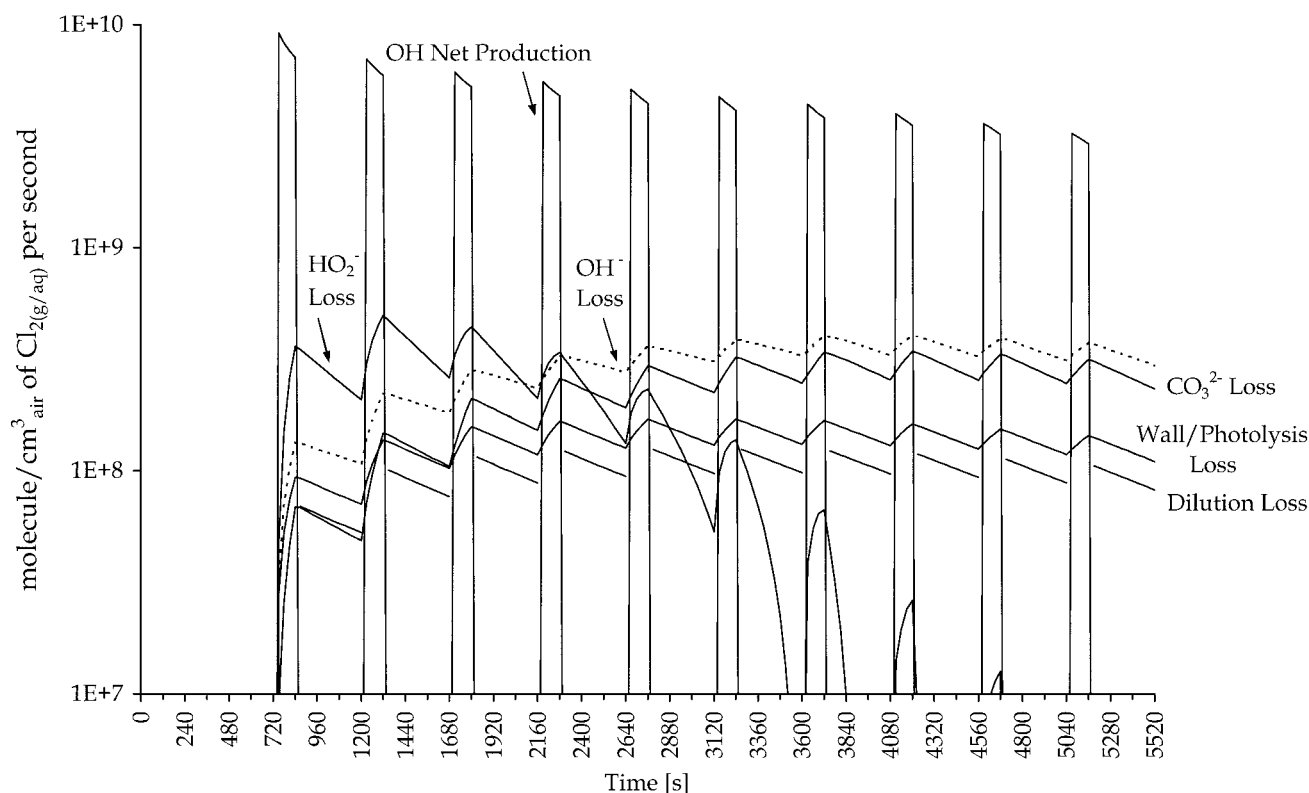


Figure 9. Rates of major Cl₂ production and loss routes.

a Cl₂ production mechanism initiated and limited by the reaction of Cl⁻ and OH[•] at aqueous interfaces.

4. Decomposition of Chlorine in Alkaline Solution

[81] The most significant sink for molecular chlorine in the experimental chamber, according to the interfacial modeling results, is its reaction with hydroxide within the droplets. This aqueous-phase reaction may of modest concern in the real atmosphere as marine aerosol pH values alkalinity most commonly remain below 8 [Keene *et al.*, 1998; Keene and Savoie, 1998, 1999; Erickson *et al.*, 1999; Fridlind and Jacobson, 2000] with a few exceptions such as within crustal dust plumes over marine regions. However, this reaction is of paramount importance in these experimental studies where model predicted pH values approach

11. Given its vital impact to these findings, a review of this reaction is warranted. In particular, the methods previously used to estimate the rate constant for the reaction of Cl₂ and OH⁻ are examined and the basis for selecting a value for use in the aqueous-phase chemical mechanism of MAGIC is presented. In addition, the other major alkaline chlorine decomposition processes, reaction with CO₃²⁻ and HO₂⁻, are also examined.

4.1. Linear Free Energy Relations

[82] In the analyses of the reaction of chlorine with hydroxide, and chlorine hydrolysis in general, two topics commonly arise: the application of linear free energy relations and the pursuit of evidence regarding the composition of the intermediate species. Linear free energy relations include the Bronsted–Pedersen relations for acid or

Table 9. Sensitivity of Peak and Final Cl₂ Concentrations Within Limits of Key Reaction Rate Constants

k_{1134}	k_{1135}	k_{1137}	k_{1215}	k_{1217}	$[\text{Cl}_2]_{\text{peak}}$	$[\text{Cl}_2]_{\text{final}}$
M ⁻¹ s ⁻¹	M ⁻¹ s ⁻¹	M ⁻¹ s ⁻¹	M ⁻¹ s ⁻¹	M ⁻¹ s ⁻¹	molecule cm ⁻³	molecule cm ⁻³
1.0 × 10⁸	2.0 × 10⁻³	1.1 × 10⁸	1.4 × 10⁵	0.012	1.50 × 10 ¹²	1.26 × 10 ¹²
1.0 × 10 ⁶	2.0 × 10 ⁻⁵				1.68 × 10 ¹²	1.51 × 10 ¹²
1.0 × 10 ⁷	2.0 × 10 ⁻⁴				1.66 × 10 ¹²	1.48 × 10 ¹²
1.0 × 10 ⁹	2.0 × 10 ⁻²				1.02 × 10 ¹²	6.85 × 10 ¹¹
1.0 × 10 ¹⁰	2.0 × 10 ⁻¹				5.05 × 10 ¹¹	2.60 × 10 ¹¹
		1.1 × 10 ⁶			1.60 × 10 ¹²	1.26 × 10 ¹²
		1.1 × 10 ⁷			1.62 × 10 ¹²	1.26 × 10 ¹²
		1.1 × 10 ⁹			1.27 × 10 ¹²	1.25 × 10 ¹²
			1.4 × 10 ³	1.2 × 10 ⁻⁴	1.62 × 10 ¹²	1.44 × 10 ¹²
			1.4 × 10 ⁴	1.2 × 10 ⁻³	1.64 × 10 ¹²	1.47 × 10 ¹²
			1.4 × 10 ⁶	0.12	1.13 × 10 ¹²	7.84 × 10 ¹¹

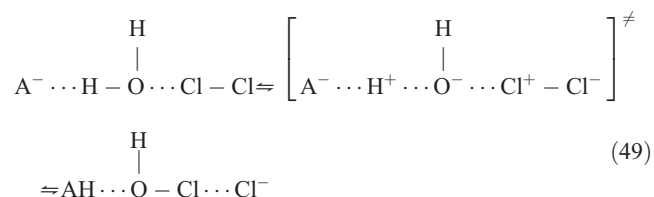
Bold values on first line are as listed on Table 5. Only values changed in the sensitivity runs are listed on other lines.

Table 10. Summary of Bronsted–Pedersen Relationship Parameters for Base-Assisted Cl₂ Hydrolysis (Cl₂ + H₂O + A[−] ⇌ HOCl + Cl[−] + HA)^a

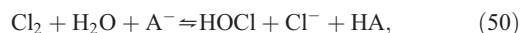
HA	A [−]	<i>p</i>	<i>q</i>	p <i>K_a</i>	<i>k_−</i> , M ^{−1} s ^{−1}	log(<i>K_a</i> <i>q/p</i>)	log(<i>k_−</i> / <i>q</i>)	Residuals fit with H ₃ O ⁺ ^b	Residuals fit w/o H ₃ O ⁺ ^c	Comments
H ₃ O ⁺	H ₂ O	3	2	−1.74	1.6 × 10 ^{−1}	1.56	−1.10	0.084		
HSO ₄ [−]	SO ₄ ^{2−}	1	4	1.27	3.2 × 10 ¹	−0.67	0.90	0.166	0.005	
CHCl ₂ COOH	CHCl ₂ COO [−]	1	2	1.30	2.7 × 10 ¹	−1.00	1.13	0.198	0.024	d
H ₃ PO ₄	H ₂ PO ₄ [−]	3	2	1.75	4.5 × 10 ¹	−1.93	1.35	0.015	0.001	
CH ₃ ClCOOH	CH ₃ ClCOO [−]	1	2	2.81	9.7 × 10 ¹	−2.51	1.69	0.014	0.002	
HCOOH	HCOO [−]	1	2	3.72	1.2 × 10 ²	−3.42	1.78	0.104	0.068	
CH ₃ COOH	CH ₃ COO [−]	1	2	4.76	9.4 × 10 ²	−4.46	2.67	0.001	0.032	
								ΣRes. = 0.582	ΣRes. = 0.131	
								<i>r</i> ² = 0.932	<i>r</i> ² = 0.937	e
								χ ² = 0.116	χ ² = 0.033	e
Extrapolated points										
CO ₂ •H ₂ O	HCO ₃ [−]	2	2	6.08	3.2 × 10 ³	−6.08	3.21			f
HCO ₃ [−]	CO ₃ ^{2−}	1	3	9.90	1.4 × 10 ⁵	−9.42	4.67			f
H ₂ O ₂	HO ₂ [−]	2	3	11.60	1.1 × 10 ⁶	−11.43	5.55			f,g
H ₂ O	OH [−]	2	3	15.74	6.9 × 10 ⁷	−15.56	7.36			f

^aData obtained from the work of Wang and Margerum [1994] and references therein at *T* = 15°C and *I* = 0.5 M.^bβ = 0.583, *G_B* = 0.541.^cβ = 0.483, *G_B* = 0.106.^dObtained from the work of Lifshitz and Perlmutter-Hayman [1962] and corrected for temperature and ionic strength effects. Point not used in derivation of fits (thus, results are comparable to fit by Wang and Margerum [1994]) but included in *r*-squared and chi-squared calculations.^e*r*² = 0.953, χ² = 0.096 (with H₃O⁺) and *r*² = 0.937, χ² = 0.036 (without H₃O⁺) if CHCl₂COOH point omitted from statistical calculations.^f*k_−* values calculated using Bronsted–Pedersen fit without H₃O⁺.^gDirect reaction with HO₂[−] to form HOOC[−] and Cl[−] overwhelms potential contribution via base-assisted mechanism (see section 4.2).

base catalyzed reactions [Benson, 1960; Bell, 1973] and the Swain–Scott relation [Swain and Scott, 1953]. In the case of chlorine hydrolysis, studies have shown clear evidence for a base-assisted forward reaction and acid-assisted reverse reaction [Lifshitz and Perlmutter-Hayman, 1961, 1962; Wang and Margerum, 1994]. Wang and Margerum [1994] present a general hydrolysis mechanism consisting of Cl⁺ and proton transfers between the reactants and the transition state, as shown by



Whereas the nature of the actual intermediate(s) formed during these reactions is still a subject of much debate, the rate constants for the overall reactions,



are found proportional to the acid–base strength of the catalyst, in agreement with the Bronsted–Pedersen equations,

$$\log \left(\frac{k_-}{q} \right) = \log G_B - \beta \log \left(\frac{K_a q}{p} \right) \quad (51)$$

$$\log \left(\frac{k_-}{p} \right) = \log G_A + \alpha \log \left(\frac{K_a q}{p} \right) \quad (52)$$

where *K_a* is the acid dissociation constant of HA, *p* is the number of dissociable equally bound protons (equivalent protons) on HA, and *q* is the number of positions that can

attach a proton (equivalent basic sites) on the conjugate base A[−]. The constants α and β are measures of the amount of free energy change of ionization attributed to the formation of the transition state and theoretically add up to unity, while *G_A* and *G_B* are proportionality constants. Table 10 lists the literature values of *k_−*, and for several acid/base pairs. Figure 10 illustrates the Bronsted–Pedersen relationship as a plot of log(*k_−*/*q*) versus log(*K_aq/p*); the slope of the line and intercept are given by β and *G_B*, respectively. (A similar line for log(*k_−*/*q*) versus log(*K_aq/p*), where the slope and intercept are given by α and *G_B*, is not shown in order to minimize clutter.) Included in Table 10 is the rate constraint for the hydrolysis of chlorine. In this case, reaction (50) is written so that the acid/base pair is taken as the hydronium ion, H₃O⁺, and water,



This point has been used by the previously cited researchers to extrapolate the value for the case of the acid/base pair of water and hydroxide, i.e.,



However, the validity of the HA/A[−] pairs H₃O⁺/H₂O and H₂O/OH[−] in the Bronsted–Pedersen relation is debatable. (The argument applies specifically for the relation as given by equations (51) and (52).) Significant departures from the linear free energy relation are not uncommon, with deviations near three orders of magnitude reported for other reactions of halogens in water [Nagy et al., 1988; Kumar and Margerum, 1987; Beckwith et al., 1996].

[83] A full discussion on the legitimacy of the species H₃O⁺, H₂O and OH[−] in the Bronsted–Pedersen relation is provided elsewhere [Bell, 1973, and references therein]. In brief, the linear free energy relation has been shown valid for catalysts of similar chemical type and charge. Some

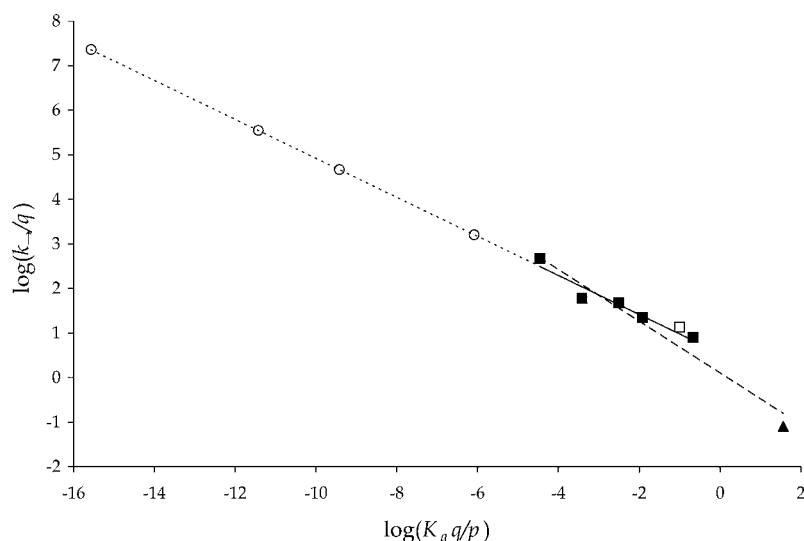


Figure 10. Bronsted–Pedersen chart for the reaction $\text{Cl}_2 + \text{H}_2\text{O} + \text{A}^- \rightleftharpoons \text{HOCl} + \text{Cl}^- + \text{HA}$ (see Table 10 for details): ■: experimental points used in original and alternate fit; □: experimental dichloroacetate point not used in either fit; ○: extrapolated points using alternate fit; —: alternate fit; - - -: alternate extrapolation; ▲: $\text{H}_3\text{O}^+/\text{H}_2\text{O}$ point used in original fit; - - -: original fit by Wang and Margerum [1994].

curvature to the relation has been found when considering dissimilar acids and bases, always showing a tendency of lowering slope (smaller α or β values) with increasing rate constants (larger pK values). The curvature becomes more apparent by inclusion of H_3O^+ , H_2O and OH^- ; these species are of different chemical type and frequently of different charge than the other acids or bases employed in catalysis studies. As a result, the hydroxide ion and (to a lesser extent) the hydronium ion have been found to frequently exhibit a catalytic effect several orders of magnitude lower than that predicted by the Bronsted–Pedersen relation.

[84] In view of the available evidence, it is entirely conceivable that inclusion of H_2O , a neutral species and the solvent, in the analysis of halogen hydrolysis assisted by negatively charged bases could lead to erroneous predictions by Bronsted–Pedersen extrapolations to hydroxide-assisted reactions. In the case of Cl_2 hydrolysis, a detailed inspection shows that two data sets, one including and another excluding the $\text{H}_3\text{O}^+/\text{H}_2\text{O}$ point, yield similarly high coefficients of linear correlation (r^2). However, a better fit ($\sim 3\times$) to the data—as measured by the normalized chi-square merit function (χ^2) with all points equally weighted—is obtained by neglecting the data point for $\text{H}_3\text{O}^+/\text{H}_2\text{O}$ (see Table 10 and Figure 10). Extrapolating the better fit to reaction with OH^- , the rate constant is predicted at $6.9 \times 10^7 \text{ M}^{-1} \text{ s}^{-1}$ at 298 K and $I = 0.5 \text{ M}$. The extrapolation overestimates the rate constant for reaction with H_2O by a factor of ~ 9 , as expected from previous arguments.

[85] Several studies of the specific and acid-assisted attack on the halogen of hypohalous acids by nucleophilic reagents [Nagy *et al.*, 1988; Gerritsen and Margerum, 1990; Wang and Margerum, 1994; Beckwith *et al.*, 1996], in particular halide ions, have shown that the rate constants for these reactions appear to follow the Swain–Scott relationship. However, most analyses using this linear free energy relation are based on a limited amount of experimental points (usually three) corresponding to permutations

of studies of HOCl , HOBr and HOI reactions with Cl^- , Br^- and I^- . In the specific case of the reaction of HOCl with nucleophilic anions (equivalent to the acid-assisted reverse hydrolysis reaction with $\text{HA} = \text{H}_2\text{O}$) [Gerritsen and Margerum, 1990], the other nucleophiles considered, CN^- and SO_3^{2-} , have nucleophilicities so similar to Cl^- (and correspondingly similar rate constants) that an indisputable preference between the previously discussed Bronsted–Pedersen extrapolations for the reaction of Cl_2 with hydroxide cannot be established from Swain–Scott analysis of the reverse reaction.

[86] In essence, the determination of the mechanisms of chlorine hydrolysis appears tantamount to the explanation of the transition state(s). Several intermediates have been explored in order to explain the hydrolysis of chlorine in aqueous solution including Cl_2OH^- , H_2OCl_2 , and H_2OCl^+ [Bell and Gelles, 1951; Anbar and Taube, 1958; Anbar, *et al.*, 1959; Eigen and Kustin, 1962; Swain and Crist, 1971; Held *et al.*, 1978; Wang and Margerum, 1994; Dahl and Roeggen, 1996; Donaldson *et al.*, 1997; Liu *et al.*, 1999]. However, a molecular-scale analysis of chlorine hydrolysis is beyond the scope of this paper.

4.2. Reaction of Dissolved Cl_2 with OH^- : Rate Constant Determination

[87] The reaction between Cl_2 and OH^- has been studied extensively, due in particular to its commercial applications in bleaching and disinfection processes. Early calculations of the rate constant of the $\text{Cl}_2 + \text{OH}^-$ reaction as $8.3 \times 10^{12} \text{ M}^{-1} \text{ s}^{-1}$ [Morris, 1946] were fraught with error [Anbar and Taube, 1958], yielding a value that exceeds even the rate constant of proton transfer reactions ($\sim 10^{11} \text{ M}^{-1} \text{ s}^{-1}$), e.g., the reaction between H_3O^+ with OH^- or SO_4^{2-} [Benson, 1960]. Progress was made when Lifshitz and Perlmuter-Hayman [1960, 1961] studied the rate of chlorine hydrolysis in pure water and in the presence of an acetate buffer at 282.5 K. Their initial studies estimated the rate constant for the reaction of Cl_2 with hydroxide at $9 \times 10^{10} \text{ M}^{-1} \text{ s}^{-1}$, but

Table 11. Summary of Reported Rate Constants for the Reaction $\text{Cl}_2 + \text{OH}^- \rightleftharpoons \text{HOCl} + \text{Cl}^-$ ^a

Year	Method	$k_{\rightarrow}, \text{M}^{-1} \text{s}^{-1}$	$\log(k_{\rightarrow})$	Reference
1946	Estimate	8.3×10^{12}	12.92	Morris [1946]
1961	Estimate	9.0×10^{10}	10.95	Lifshitz and Perlmutter-Hayman [1961]
1962a	Estimate	1.0×10^{10}	10.0	Eigen and Kustin [1962]
1962b	Bronsted–Pedersen Extrapolation	1.0×10^9	9.0	Lifshitz and Perlmutter-Hayman [1962]
1962c	Alternative Bronsted Extrapolation	1.0×10^8	8.0	Alternate interpretation of above data
1962d	Measurement	1.0×10^6	6.0	Spalding [1962]
1981	Measurement	2.7×10^7	7.43	Sandall et al. [1981]
1994a	Bronsted–Pedersen Extrapolation	4.1×10^9	9.61	Wang and Margerum [1994]
1994b	Alternative Bronsted–Pedersen Extrapolation	6.9×10^7	7.84	Alternate interpretation of above data, this work
1996	Measurement	1.6×10^9	9.19	Ashour et al. [1996]
2001	Measurement	5.0×10^8	8.70	Gershenson et al. (in preparation)

^aMeasurements and extrapolations are taken at different temperatures (in the range 273–298 K) and ionic strengths (in the range $I < 1$ M).

this was modified promptly [Lifshitz and Perlmutter-Hayman, 1962] with a value of $10^9 \text{ M}^{-1} \text{ s}^{-1}$. The lower rate constant value was derived from extrapolation of the Bronsted–Pedersen law to OH^- after observing evidence for a general base-assisted Cl_2 hydrolysis mechanism. The Bronsted–Pedersen parameters used in the extrapolation were obtained using only “the two most reliable” points, $\text{CH}_3\text{COO}^-/\text{CH}_3\text{COOH}$ and $\text{CH}_2\text{ClCOO}^-/\text{CH}_2\text{ClCOOH}$. If all the points except for $\text{H}_2\text{O}/\text{H}_3\text{O}^+$ (questionable in the Bronsted–Pedersen relation, vide supra) and the point for $\text{H}_2\text{PO}_4^-/\text{H}_3\text{PO}_4$ (unclear due to possible parallel reactions involving $\text{HPO}_4^-/\text{H}_2\text{PO}_4^-$) are considered, the Bronsted–Pedersen extrapolation yields a value of $\sim 10^8 \text{ M}^{-1} \text{ s}^{-1}$. Lifshitz and Perlmutter-Hayman admit that “the extrapolation [to OH^-] is so far that it cannot be expected to be accurate. It would seem desirable to confirm the value of k_{OH^-} by additional evidence.”

[88] Eigen and Kustin [1962] studied the kinetics of halogen hydrolysis by temperature-jump relaxation methods and estimated the rate of $\text{Cl}_2 + \text{OH}^-$ reaction at $10^{10} \text{ M}^{-1} \text{ s}^{-1}$. This value has been widely cited although it was obtained from a mathematical analysis of experiments performed in the narrow and acidic pH range between 1 and 2.2 and involved important assumptions regarding the intermediate structures formed during hydrolysis.

[89] By analyzing chlorine absorption data in aqueous media (pH range of 1–14.2) using penetration theory, Spalding [1962] was able to define four distinct pH ranges (using H_2SO_4 and NaOH as buffers) that characterized Cl_2 absorption kinetics due to different hydrolysis effects. The reaction of Cl_2 with OH^- progressively governed the absorption rate as pH values increased above 10.5 until ultimately controlling the absorption rate at $\text{pH} \geq 12.5$. Spalding estimated a rate constant for the $\text{Cl}_2 + \text{OH}^-$ reaction at $10^6 \text{ M}^{-1} \text{ s}^{-1}$.

[90] A lack of consensus became apparent when Cl_2 absorption data obtained from a roller-drum reactor, analyzed using a surface renewal model, yielded a rate constant of $2.7 \times 10^7 \text{ M}^{-1} \text{ s}^{-1}$ at 273 K for the chlorine hydroxide reaction [Sandall et al., 1981]. No conclusion can be drawn about the room temperature rate constant as the authors did not explore this.

[91] Wang and Margerum [1994] examined the kinetics of reversible chlorine hydrolysis, reconciling disagreement in previously reported data by providing the temperature

and ionic strength dependence of the forward and reverse reactions as well as the equilibrium constant. In addition, Wang and Margerum investigated general base assisted mechanisms for Cl_2 hydrolysis, augmenting the data of Lifshitz and Perlmutter-Hayman [1961, 1962] by evaluating the chlorine hydrolysis rate in the presence of sulfate and phosphate buffers using a stopped-flow method with spectrophotometric observation. The rate constants obtained were combined with those of Lifshitz and Perlmutter-Hayman for monochloroacetate, formate and acetate (after correcting for temperature and ionic strength effects) in order to obtain the parameters for the Bronsted–Pedersen equations. In doing so, the authors included the point for $\text{H}_2\text{O}/\text{H}_3\text{O}^+$ and extrapolated their “best fit” to the point for $\text{OH}^-/\text{H}_2\text{O}$ obtaining a value of $4.6 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$. Furthermore, the authors match up their result to that of Eigen and Kustin [1962] by proposing that an additional direct reaction between Cl_2 and OH^- with a rate constant of $6.3 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$ would also contribute to the formation of a Cl_2OH^- intermediate, therefore setting the overall rate constant for $\text{Cl}_2 + \text{OH}^-$ near $10^{10} \text{ M}^{-1} \text{ s}^{-1}$.

[92] Ashour et al. [1996] measured the rate of Cl_2 absorption into aqueous sodium hydroxide (0.1 M NaOH) solutions in the temperature range of 293–312 K. Extrapolating the resulting temperature dependent expression, the rate constant for the reaction of $\text{Cl}_2 + \text{OH}^-$ is calculated at $1.3 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$ at 288 K ($2.9 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$ at 298 K), below that predicted by the Bronsted–Pedersen extrapolation using $\text{H}_2\text{O}/\text{H}_3\text{O}^+$ but nevertheless above that predicted without using hydronium in the linear free energy relation (see Table 10 and Figure 10).

[93] In the previous section, an alternate Bronsted–Pedersen extrapolation predicted a rate constant for the reaction of Cl_2 with OH^- of $6.9 \times 10^7 \text{ M}^{-1} \text{ s}^{-1}$ at 298 K and $I = 0.5$ M. During the preparation of this manuscript, a recent measurement of the $\text{Cl}_2 + \text{OH}^-$ rate constant at 298 K using a bubble column apparatus [Swartz et al., 1997] of $5 (\pm 2) \times 10^8 \text{ M}^{-1} \text{ s}^{-1}$ was communicated to the authors (M. Gershenson et al., Rate constant for the $\text{Cl}_{2(aq)} + \text{OH}^-$ reaction, manuscript in preparation). Table 11 summarizes reported rate constants for the reaction. These measurements and estimates illustrate that at present there is still a lack of consensus regarding one definitive and reproducible value. A rate constant of $1 \times 10^8 \text{ M}^{-1} \text{ s}^{-1}$ (at

infinite dilution and 298 K), near the logarithmic center of the lowest and highest reported rate constants, has been chosen for MAGIC as the best compromise of the data scatter.

[94] The selection of a dissimilar value from the most recent experimental report does not imply any bias save for the indication that further studies are needed in order to reach a generally reproducible value. Model simulations adequately predict Cl₂ concentrations after the first three photolysis periods (when the aqueous-phase reaction of Cl₂ with HO₂⁻ predominates over the reaction with OH⁻) applying the value of $5 \times 10^8 \text{ M}^{-1} \text{ s}^{-1}$ for the chlorine hydroxide reaction rate constant and using $\gamma'_s = 2$. However, the simulations increasingly underestimate Cl₂ levels as time progress, with model calculations approaching 1/2 of the experimental measurement at the end of the simulations. Clearly, further experimental work is needed not only to better characterize the interfacial mechanism but also to better determine the rates of Cl₂ reactions in alkaline solution.

[95] The reported values in Table 11 should be viewed with the recognition that many rate constants were determined at different temperatures and ionic strengths. The inconsistency in the studies do not allow for a clear deduction of temperature effects, though it is noted that *Ashour et al.* [1996] do report a weak temperature. However, using the temperature dependence of the chlorine hydrolysis equilibrium constant, the water equilibrium product, K_w [Whitfield, 1975],

$$\log K_w = -4470.99/T + 6.0875 - 0.01706 T \text{ (at infinite dilution)} \quad (55)$$

and activity coefficient estimates, the following equilibrium expression

$$K = [\text{HOCl}][\text{Cl}^-]/[\text{Cl}_2][\text{OH}^-] \quad (56)$$

can be determined, with a value of 9×10^{10} at 273 K and a value of 4×10^{10} at 298 K (both at $I = 0.5 \text{ M}$). At infinite dilution, the equilibrium constant increases by a factor of 1.3 while at higher ionic strengths short-range molecular interactions become important requiring individual determination of activity coefficients.

4.3. Reaction of Dissolved Cl₂ with CO₃²⁻: Rate Constant Determination

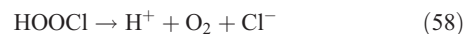
[96] *Ashour et al.* [1996] determined an expression for the rate constant of $\text{Cl}_2 + \text{HCO}_3^- + \text{H}_2\text{O}$ in the temperature range of 293–312 K by studying potassium bicarbonate solutions (0.1 to 0.4 M initial concentration) at pH values near 8.5. Their experimental conditions were such that $[\text{CO}_3^{2-}]/[\text{HCO}_3^-] \approx 0.011$. The authors assumed in their penetration model that the rate constant for $\text{Cl}_2 + \text{CO}_3^{2-} + \text{H}_2\text{O}$ equals that of $\text{Cl}_2 + \text{HCO}_3^- + \text{H}_2\text{O}$, an assumption that both extrapolations of the Bronsted–Pedersen relation oppose. Extrapolating their results to 288 K and $I = 0.5 \text{ M}$, a value for $k_{\text{HCO}_3^-}$ (and thus also for $k_{\text{CO}_3^{2-}}$) of $2.6 \times 10^3 \text{ M}^{-1} \text{ s}^{-1}$ is computed ($3.7 \times 10^3 \text{ M}^{-1} \text{ s}^{-1}$ at 298 K). In comparison, at the same temperature and ionic strength, the Bronsted–Pedersen interpolation without $\text{H}_2\text{O}/\text{H}_3\text{O}^+$ yields $k_{\text{HCO}_3^-} = 3.2 \times 10^3 \text{ M}^{-1} \text{ s}^{-1}$ and $k_{\text{CO}_3^{2-}} = 1.4 \times 10^5 \text{ M}^{-1}$

s^{-1} while applying the free energy relation with $\text{H}_2\text{O}/\text{H}_3\text{O}^+$ gives $k_{\text{HCO}_3^-} = 9.0 \times 10^3 \text{ M}^{-1} \text{ s}^{-1}$ and $k_{\text{CO}_3^{2-}} = 1.2 \times 10^6 \text{ M}^{-1} \text{ s}^{-1}$. Bearing in mind the apparently erroneous assumption, $k_{\text{HCO}_3^-} = k_{\text{CO}_3^{2-}}$, the rate constant calculated by Ashour et al. is more consistent with the hydronium-less (alternate) Bronsted–Pedersen extrapolation. In view of the uncertainties, the alternate extrapolated values, even if calculated at 288 K, are included in the model.

4.4 Reactions of Dissolved Cl₂/HOCl with H₂O₂/HO₂⁻: Rate Constant Determination

[97] In the model simulations, concentrations of hydrogen peroxide reach appreciable levels. In lieu of the preceding discussion, the reaction of Cl₂ with HO₂⁻ could be expected to follow a base-assisted mechanism, but there is evidence leading to another predominant reaction pathway.

[98] *Connick* [1947] hypothesized that the mechanism for the reaction of chlorine in hydrogen peroxide solution could be explained by the decomposition of a chloroperoxide species, HOOC₂Cl, formed from the direct reaction of Cl₂ with H₂O₂,



with a rate law in the form:

$$-d[\text{H}_2\text{O}_2]/dt = k_a[\text{H}_2\text{O}_2][\text{Cl}_2]/(k_b/k_c)[\text{H}^+][\text{Cl}^-] + 1 \quad (59)$$

with $k_a = 183.3 \text{ M}^{-1} \text{ s}^{-1}$ and $(k_b/k_c) = 2.27 \text{ M}^{-2}$. This expression matched data from earlier studies [Makower and Bray, 1933] at high HCl concentrations and remains valid down to $[\text{H}^+][\text{Cl}^-]$ products of 10^{-5} M^2 , i.e., hydrochloric acid concentrations $>0.003 \text{ M}$. *Davies and Kustin* [1973] obtained $k_a = 192 \text{ M}^{-1} \text{ s}^{-1}$ and $k_b/k_c = 1.6 \text{ M}^{-2}$ from a least squares analysis of Connick's data, a small difference due possibly to different values of the Cl₂ hydrolysis equilibrium constant.

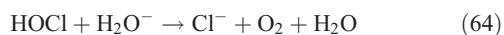
[99] At $[\text{H}^+][\text{Cl}^-]$ products below 10^{-8} M^2 (hydrochloric acid concentrations below 10^{-4} M , alternatively), HOCl becomes the main chlorine species in solution and the rate of H₂O₂ decay observed by Connick is much greater than could be explained by Expression 59. The rate law for the new reaction path detected by Connick is given by

$$-d[\text{H}_2\text{O}_2]/dt = k[\text{H}_2\text{O}_2][\text{HOCl}]/[\text{H}^+] \quad (60)$$

with $k \approx 10^{-4} \text{ M}^{-1} \text{ s}^{-1}$. Connick also proposed that the decay of HOCl in H₂O₂ solution could be better explained by a hypochlorite pathway described by

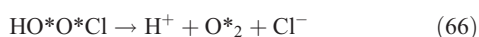


instead of a HO₂⁻ pathway

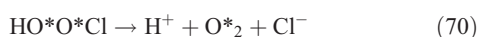
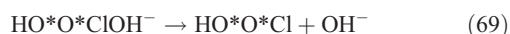


on grounds that the Arrhenius constant of the latter pathway appeared “too great.” Connick also observed evidence for a third reaction pathway in the region where $10^{-5} > [\text{H}^+][\text{Cl}^-] > 10^{-8}$ but was unable to provide a corresponding rate law.

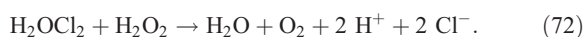
[100] *Cahill and Taube* [1952] performed experiments on the oxidation, reduction and catalytic decomposition of H₂O₂ using heavy oxygen, O¹⁸, as a tracer. They found that the O₂ liberated during the oxidation of H₂O₂ was derived “cleanly” from H₂O₂. The H–O bonds rather than the O–O bond were cleaved, leading to the conclusion that oxidation proceeded by removal of electrons from H₂O₂, thus strengthening the O–O bond. Cahill and Taube’s finding supported the high acid range mechanism of Connick and the formation of an HOOC₂ intermediate,



but favored the alternative mechanism proposed for the low acid concentration range. An activated complex, HOOC₂OH[−], was proposed, consistent with a mechanism given by



[101] Further studies showed that the molecular oxygen formed during the decomposition of hydrogen peroxide and sodium hypochlorite to be in the electronically excited ¹Δ_g state, recognized by its distinctive dimol chemiluminescence at 633.4 and 762.0 nm [*Khan and Kasha*, 1963, 1970]. Chemiluminescence measurements identified either the HOCl/HO₂[−] or the ClO[−]/H₂O₂ reactions as the source of the singlet oxygen, but were only able to discard likelihood of a mechanism involving hydride ion transfer from HO₂[−] to ClO[−] [*Kajiwara and Kearns*, 1973]. *Held et al.* [1978] determined that the least energetically demanding path would be attack by the strongly nucleophilic HO₂[−] ion on the electropositive chlorine atom of HOCl to form the HOOC₂ intermediate, supporting the mechanism of *Cahill and Taube* [1958], with a rate constant of $4.4 \times 10^7 \text{ M}^{-1} \text{ s}^{-1}$. *Held et al.* also found evidence for an “intermediate pH” reaction pathway for reactions of Cl₂/HOCl in solutions of hydrogen peroxide. Intriguingly, the rate law observed in this range, $10^{-5} > [\text{H}^+][\text{Cl}^-] > 10^{-8}$, coincided with the rate law and kinetics measured for the reverse chlorine hydrolysis reaction, $k[\text{HOCl}][\text{Cl}^-][\text{H}^+]$. The authors provide a detail analysis in support of a reaction between a Cl₂ hydrolysis intermediate, H₂OCl₂, and hydrogen peroxide



[102] This reaction could not be included in MAGIC primarily due to the inability of the authors to determine elementary kinetic data in accordance with the proposed mechanism and the apparent absence of later research on the reaction. However, the model evaluations of MAGIC predict that the aerosol particles will become highly alkaline thereby lessening the error due to this omission, and heightening the importance of reactions with HO₂[−].

[103] Although the HOOC₂ molecule has not yet been observed, the molecule has been suggested as an intermediate in gas-phase reactions and heterogeneous acid-catalyzed ice surface reactions [*Wofsy et al.*, 1988; *Lee and Rendell*, 1993; *De Haan and Birks*, 1997] and in aqueous-phase reactions of Cl₂ in basic hydrogen peroxide solutions [*Storch et al.*, 1983]. The rate constant for the reaction of Cl₂ with HO₂[−] was determined from measurements of chlorine absorption into laminar liquid jets with initial concentration [HO₂[−]] = 4 M [*Ruiz-Ibanez and Sandall*, 1991; *Davis et al.*, 1992]. A base-assisted chlorine hydrolysis mechanism cannot account for the observed singlet delta oxygen formation, O₂(¹Δ_g), observed during this reaction—the excited oxygen formation is more consistent with the formation of the HOOC₂ intermediate. *Davis et al.* [1992] reevaluated the earlier analysis of *Ruiz-Ibanez and Sandall* [1991] accounting for the reduction of solubility and diffusivity of Cl₂ ascribed to ion effects in the liquid phase of basic hydrogen peroxide solutions and obtained a value for the rate constant (using their temperature dependent expression) of $1.8 \times 10^7 \text{ M}^{-1} \text{ s}^{-1}$ at 288 K and a rate constant of $1.1 \times 10^8 \text{ M}^{-1} \text{ s}^{-1}$ extrapolating to 298 K. The possibility of a reaction between Cl₂ and OH[−] affecting the results of *Davis et al.*, can be eliminated using the recommended rate constant (section 4.2) and a calculated value of 10^{-3} M for [OH[−]] in the experiments. Furthermore, any contribution from a secondary base-assisted mechanism is also neglected since the resultant rate constant, using the Bronsted–Pedersen relation, is orders of magnitude less than that for the reaction leading to the HOOC₂ intermediate.

[104] Finally, the reaction rate of H₂O₂ with chloride ions [*Mohammad and Liefbhafsky*, 1934; *Hansen and Espenson*, 1995] is too slow to affect the system modeled by MAGIC at all.

5. Conclusion and Future Applications

[105] A model of gas-phase and aqueous-phase chemistry, MAGIC, has been developed in order to explain the production of Cl₂ from the photolysis of mixtures of ozone and deliquesced sodium chloride particles in a laboratory chamber. The *Schwartz* [1986] method is used to calculate the rate of mass transfer of soluble species considering the limitations imposed by gas-phase diffusion, mass accommodation, and aqueous-phase diffusion. Activity coefficients are calculated using the *Pitzer* [1991] ion interaction approach and the kinetic salt effect is estimated using Debye–Huckel–Bronsted equations. The model shows that within this seemingly simple system, experimental and model results cannot be reconciled on the basis of currently known gas-phase and aqueous-phase chemical reactions and mass transfer parameters.

[106] As part of a multidisciplinary research effort, collaborators have shown that chloride ions reside at the

surface of deliquesced sodium chloride particles and that these ions are expected to attract reactive gases such as $\text{OH}^\bullet$ and O_3 [Knipping *et al.*, 2000]. A methodological analysis of the system strongly supports a hypothesis for a mechanism of Cl_2 formation rate-determined by the reaction of chloride ions and $\text{OH}^\bullet$ radicals at the gas-liquid interface. A kinetic expression given for the overall mechanism matches the observed experimental data well. Sensitivity studies demonstrate that the modeling of Cl_2 production is greatly affected by the uncertainties in the rate constants for decomposition of Cl_2 in alkaline solution, in particular reaction with OH^- , HO_2^- , and CO_3^{2-} , and uncertainties in the aerosol population parameters.

[107] The reactive fate of a potential $\text{OH}\cdots\text{Cl}_{\text{surface}}^\bullet$ species formed in the rate-determining interfacial reaction is still subject to further studies. For instance, it is possible that the bound complex may self-react forming Cl_2 and hydroxide ions. In an alternative, possibly parallel mechanism, the chloride-hydroxyl surface complexes may react with other surface chloride ions forming $\text{Cl}_2^\bullet$. If the $\text{Cl}_2^\bullet$ radical anion is stabilized at the interface, it could also recombine or react with other chloride-hydroxyl complexes hence releasing chlorine to the gas phase. Considering that the $\text{OH}\cdots\text{Cl}_{\text{surface}}^\bullet$ surface complex and the species formed by OH_{aq} and Cl^- in bulk solution, $\text{HOCl}^\bullet$, are most likely not the same species, molecular dynamics simulations suggest that $\text{OH}\cdots\text{Cl}_{\text{surface}}^\bullet$ is sufficiently stable in order to accumulate and allow for the proposed interfacial processes to occur at significant rates. In summary, the overall phenomenological model presented herewith is founded on the results of quantum chemical calculations and molecular dynamic simulations, evidenced by experimental observations and supported by a series of kinetic modeling simulations which fit the observations within rational tolerances.

[108] It is important to address the relevance of the interfacial mechanism to conditions representative of the true tropospheric multiphase system. There are several atmospheric conditions worthy of investigation using these results, e.g., the remote marine boundary layer, polluted coastal regions, and the polar troposphere. A significant discussion on any of these topics requires a description of the formulation, chemical mechanism, input data, results, and analysis of an extensive modeling study, i.e., a level of detail falling beyond the scope of the current manuscript. However, some preliminary observations and comments are made at the present time.

[109] Hydrocarbon and DMS measurements in the marine boundary layer have been used to infer chlorine atom concentrations or its contribution to oxidation of these compounds in the marine boundary layer [Rudolph *et al.*, 1996; Singh *et al.*, 1996a, 1996b; Wingenter *et al.*, 1996, 1999; Davis *et al.*, 1999; Ramacher *et al.*, 1999]. Estimates of chlorine atom concentrations from these studies range from less than 10^3 molecule cm^{-3} up to 10^5 molecule cm^{-3} . The two order-of-magnitude concentration range appears to be highly dependent on pH and the pollutant load of the air masses. Previous estimates of chlorine atom concentrations using the interfacial mechanism show that these results can be reasonably reproduced [Knipping *et al.*, 2000]. Inference of 24-hour average $[\text{Cl}^\bullet]$ for the remote marine boundary layer conditions yielded 780 molecule cm^{-3} [Wingenter *et al.*, 1999]; the model predicts 980 molecule cm^{-3} . Inference

of local noon instantaneous chlorine atom concentrations for relatively polluted marine air masses yielded 6.5×10^4 molecule cm^{-3} [Wingenter *et al.*, 1996]; the model estimates 3.7×10^4 molecule cm^{-3} . The contribution of the interfacial mechanism was estimated at 40% for the remote case (pH ~ 4) and at 20% (pH ~ 3.2) for the more polluted case, leading the authors to conclude that the interfacial mechanism may begin to govern Cl atom concentrations at aerosol pH values above 4.

[110] A few additional comments seem appropriate. First, in agreement with other modeling research (R. Sander, personal communication), the principal mechanism for chlorine atom production in the marine boundary layer for the above conditions is hydrochloric acid displacement (due to acidification of the particles by nitric, sulfuric, formic, acetic and methanesulfonic acid) and subsequent reaction with $\text{OH}^\bullet$ radicals. The contribution of this acid displacement-initiated mechanism decreases with increasing pH. The rate constant of the $\text{HCl} + \text{OH}^\bullet$ reaction is rather low. Thus, at pH values above 3, the Cl_2 release (and subsequent photolysis) due to the pH-independent interfacial mechanism can be on the same order or even more effective than the HCl -hydroxyl reaction with respect to the production of chlorine atoms. However, conditions that lead to acidification are commonly those that also lead to higher $\text{OH}^\bullet$ concentrations, i.e., higher pollutant loads and correspondingly higher ozone levels. Thus, even though the interfacial mechanism may dominate chlorine atom levels at higher pH, the absolute value of the concentrations may be low, particularly in pristine conditions.

[111] Second, the contribution of the interfacial mechanism to chloride deficits measured in the atmosphere [Ayers *et al.*, 1999] is minimal. Hydrochloric acid displacement remains the dominant pathway for Cl^- depletion, although there is evidence suggesting additional mechanisms [Volpe *et al.*, 1998]. In agreement with other model studies [Sander and Crutzen, 1996; Erickson *et al.*, 1999], fresh sea salt particles are found to drop in pH quickly in timescales an hour or less; the pH to which the particles initially equilibrate depends highly on the conditions being simulated.

[112] Finally, the implications of these finding may be explored by other laboratory and atmospheric models [Sander and Crutzen, 1996; Erickson *et al.*, 1999; Herrmann *et al.*, 2000; Leriche *et al.*, 2000; Moldanova and Ljungstrom, 2001]. For example, Behnke *et al.* [1999] performed experiments in Teflon bags containing ozone, pure wet NaCl aerosol and a combination of hydrocarbons. These mixtures were irradiated using a solar illuminator. Chlorine atom production rates were deduced from measurements of hydrocarbon degradation. In all experiments where the initial ozone was below 500 ppb, no appreciable chlorine atom production was inferred. These experiments differ from those described in this paper not only in O_3 concentrations but also in the addition of hydrocarbons, the photolysis of Cl_2 , and dissimilarities in many parametric data. A detailed model study of a system comparable to the Behnke *et al.* experiments may provide further insight on the behavior of the proposed interfacial process and help ascertain the reasons for not deducing appreciable $\text{Cl}^\bullet$ levels.

[113] Modeling results for conditions typical of the Southern California Air Basin shall be presented in a forthcoming

paper (Knipping and Dabdub, in preparation). In this study we shall explore the impact chlorine chemistry, including the proposed interfacial mechanism and other heterogeneous processes. We shall also compare our results with measurements of Cl_2 in other coastal regions [Spicer *et al.*, 1998] and determine whether similar values can be reproduced.

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